14.4 MOUNTAIN- SLOPE AFFORESTATION FOR VALLEY URBAN AIR-QUALITY IMPROVEMENT

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Abstract Lanzhou is one of the major cities in northwest China and the capital of Gansu Province and located at a narrow (2-8 km width), long (40-km), NW-SE oriented valley basin (elevation: 1,500- 1,600-m) with the Tibetan plateau in the west, Baita mountain (above 1,700-m elevation) in the north, and the Gaolan mountain in the Due to topographic and meteorological characteristics, Lanzhou is one of the most polluted cities in China. Meteorological conditions (low winds, stable stratification especially inversion), pollutant sources and sinks affect the air quality. Afforestation changes the mountain-valley local circulation system, destabilizes the atmosphere, and weakens the inversion. Besides, it may absorb some pollutants (sink). Lanzhou local government carried out afforestation and pollutant-source reduction (closing several heavy industrial factories) to improve the air-quality for the past two decades. Numerical model (RAMS-HYPACT) simulates the effect of afforestation on the air pollution (TSP, SO₂, NO_x ...) control.

1. Introduction

Lanzhou is located at a narrow (2-8 km width), long (40-km), northwest-southeast oriented valley basin enclosed by 1,600-m elevation with the Tibetan plateau in the west, Baita mountain (above 1,700-m elevation) in the north, and the Gaolan mountain in the south (Fig. 1a). The highest elevation in the surroundings is the top of the Gaolan mountain around 2,150-m above the sea level.

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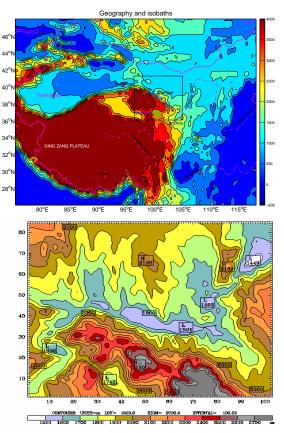


Fig. 1. Topography: (a) China, and (b) Great Lanzhou metropolitan area and surroundings.

Meteorological characteristics over the valley (great Lanzhou metropolitan) are semiarid, weak wind, thick and strong inversion, which makes the low layer atmosphere very stable, low dispersion and causes severe air pollution (Fig. 2). Lanzhou is one of the most polluted cities in China. How can air pollution be effectively controlled in valley urban area? Two possible approaches can be adopted: (1) changing meteorological conditions (destabilizing atmosphere) and (2) reducing the pollution sources.

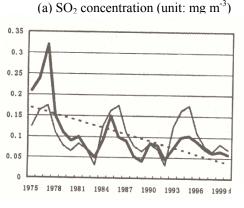


Fig. 2. LANDSAT image.

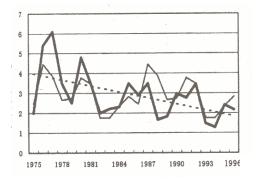
2. Air-Quality Improvement in Past Decade

Since 1980s, afforestation on the mountain slope and closure several factories that emitted large amount of pollutants were conducted. The major air pollutants (SO₂, CO, TSP, NOx) have strong seasonal variation with high values in winter and low values in summer and interannual periodicity (Wang et al. 2001). After filtering out the periodic variability, a negative trend was obtained in annual mean concentration for all pollutants (Fig. 3).

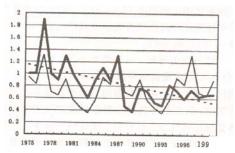
All the pollutants except NOx have maximum values in 1977. Annual mean concentration of (SO₂ CO, TSP) decreases evidently from maximum values of (0.32, 6.1, 1.9) mg/m³ in 1977 with rates of (-0.053, -0.1025, -0.0251) mg/m³ per year. Annual mean concentration of NOx has four peak years: 1980, 1985, 1990, and 1995. Negative trend in NOx concentration is much weaker (-0.0004 mg/m³ per year) than the other three pollutants.



(c) CO concentration (unit: mg m⁻³)



(c) TSP concentration (unit: mg m⁻



(d) NOx concentration (unit: mg m⁻³)

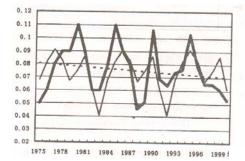


Fig. 3. Annual mean concentration (mg/m³) (a) SO₂, CO, TSP, NOx. Here, bold solid curves represent the original data, thin solid curves represent the interannual periodic variation, and dashed line represent the trend.

3. Mountain-Valley Circulation

Thermal heterogeneity of land surface can produce local circulations as strong as sea breezes (e.g., Chu 1987). Differential surface heating on the mountain slope generates local valley winds especially in winter or night (Fig. 4). The downward motion over the valley makes the atmosphere stable, and in turn weakens the diffusion of the pollutants.

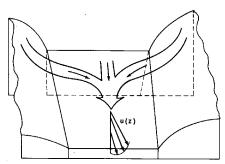


Fig. 4. Mountain-valley circulation.

Weakening this mountain-valley circulation (strong downward branch over the valley) destabilizes the atmosphere and enhances the diffusion rate. From physical point of view, reduction of surface thermal inhomogeneity will weaken this circulation. Afforestation on the mountain slope may reduce the thermal inhomogeneity and in turn improve the airquality by atmospheric destabilization. Furthermore, the forest may also absorb pollutants (as pollutant sink).

4. Atmospheric Destabilization due to Mountain Slope Afforestation

This problem is investigated using the Regional Atmospheric Modeling System (RAMS).

4.1. Model Description

RAMS is a mesoscale modeling system including advanced model physics was developed by The Colorado State University. It

is a community regional model widely used for numerical weather prediction, hydrological studies, and air quality studies. The nonhydrostatic RAMS is used in this study.

The land surface model is coupled to RAMS to describe the effect of vegetation and interactive soil moisture on the surfaceatmosphere exchange of momentum, heat, and moisture. This LSM is able to provide not only reasonable diurnal variations of surface heat fluxes as surface boundary conditions for coupled models, but also correct seasonal evolutions of soil moisture in the context of a long-term data assimilation system. Also, 1-km resolution vegetation and soil texture maps are introduced in the coupled RAMS-LSM system to help identify vegetation/water/soil characteristics at fine scales and capture the feedback of these land surface forcing. A monthly varying climatological $0.15^{\circ} \times 0.15^{\circ}$ green vegetation fraction is utilized to represent the annual control of vegetation on the surface evaporation.

LSM has one canopy layer and the following prognostic variables: soil moisture and temperature in the soil layers, water stored on the canopy, and snow stored on the ground. Four soil layers are used to capture the evolution of soil moisture and to mitigate the possible truncation error in discretization. The thickness of each soil layer from the ground surface to the bottom is 0.1, 0.3, 0.6 and 1.0 m. The precipitation is parameterized by several different schemes. Non-convective precipitation can be represented via an implicit scheme, whereby supersaturated water immediately precipitates, and an explicit scheme including prognostic equations for and rainwater. Convective water precipitation is parameterized via two cumulus convection schemes. We use the mass flux scheme, which accounts for the effects of penetrative downdrafts (Grell et al. 1994). In the numerical simulation, a flat bottom with elevation of 1460 m is assumed. This indicates that 850 hPa level is nearly at the land surface.

Twenty-three vertical levels are used with 10 hPa at the top of the atmosphere. See the RAMS website:

http://www.rams.atmos.colostate.edu for more information

4.2. Triple-Nested Grid System

A triple-nested grid systems (Fig. 5) with the same center located at (35.1°N, 103.8°E) is used in this study. The first system (large) extends 720 km in the north-south direction and 540 km in the east-west direction with the grid spacing of 9 km. The second system (medium) extends 270 km in the north-south direction and 216 km in the east-west direction with the grid spacing of 3 km. The third system (small) extends 102 km in the north-south direction and 84 km in the east-west direction with the grid spacing of 1 km.

Lanzhou is located in the smallest box. Roy and Avissar (2000) characterize the convective boundary layer (CBL) over domains with meso-gamma-scale (2-20 km) heterogeneity and find two typical length-scales of the processes when the length-scale of the heterogeneity exceeds 5-10 km: (a) 1.5 times the CBL height for turbulent thermals and (b) heterogeneity scale for organized eddies. Only the simulation in the smallest box is used for the analysis.

4.3. Boundary and Initial Conditions

At the surface, we use USGS vegetation 25-catogory with type-1 for urban/built-up land, and type-4 for mixed dry/irrigational plants (afforestation). The NCEP data along the lateral boundary (every 6 hours) of the largest box from December 1 to 31, 2000 are taken as the open boundary conditionOne way nesting is used for the triple-nested grid system. The larger model provides the lateral boundary conditions for the smaller model using a 5 point-buffer zone.

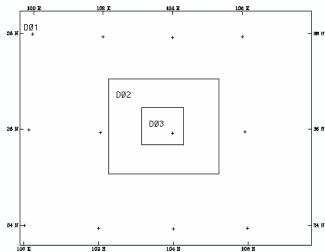


Fig. 5. Triple nesting grid system.

The NCEP reanalysis data on December 1, 2000 are taken as the initial condition. The time step is 60 s for the first grid system, 30 s for the second grid system, and 10 s for the third grid system.

4.3. Numerical Experiments

Two numerical experiments are conducted: (1) with mountain-slope afforestation, and (2) mountain-slope without afforestation. The difference between the two is the land surface. The soil type on the mountain-slope is 4 for Exp-1 (Fig. 6) and 1 for Exp-2. Everything else is kept the same for the two experiments. Model difference (Exp-1 minus Exp-2) is analyzed especially the stratification and velocity field. Afforestation decreases the downward motion over the valley (i.e., reduction of the mountainvallev circulation) and stratification (destabilization). The maximum reduction of the stratification is over the valley (-8°K/km). Such conditions favor the dispersion of valley urban air pollutants and improvement of the air-quality.

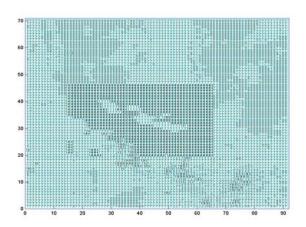


Fig. 6. Surface condition for mountain-slope afforestation with the soil type-4. The soil type-4 is replaced by type-1 for experiment without mountain-slopeaforestation.

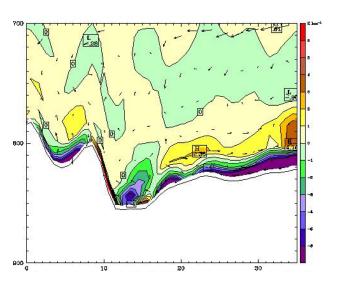


Fig. 7. Latitudinal cross-section along 103.8°E of differential (Exp-1 minus Exp-2) stratification (color contour) and (v, w) velocity vectors.

5. Air Pollution Modeling

5.1. Model Description

The Hybrid Particle and Transport (HYPACT) model developed by the Mission Research Corporation (Walko et al., 2001) is used to predict the dispersion of air pollutants in 3-D, mesoscale, time dependent wind and turbulence field. HYPACT allows assessment of

the impact of one or multiple sources emitted into highly complex local weather regimes, including mountain-valley and complex terrain flows. In this study, the modeling flow chart is shown in Fig. 8. The NCEP data is used to initialize the mesoscale model (RAMS), which provide the velocity field as input to the dispersion model (HYPACT).

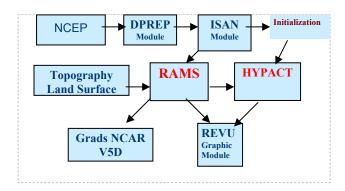


Fig. 8. RAMS/HYPACT modeling flow chart.

5.2. Pollutant Sources

Emission rates from the ground pollutant sources (industrial and residential) were measured such as SO₂, NOx, CO, TSP, etc.. RAMS/HYPACT is integrated from Dec 1, 2000 with the observational pollutant sources to Dec 31, 2000.

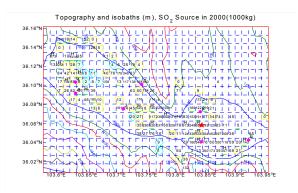


Fig. 9. SO₂ sources.

5.3. Model Verification

During the prediction period (Dec 1-31, 2000), eight observations of SO_2 were conducted on Dec 25, 2000. Except station-1, the predicted and observed SO2 concentrations agree with each other quite well.

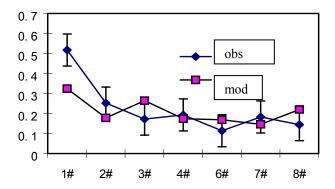
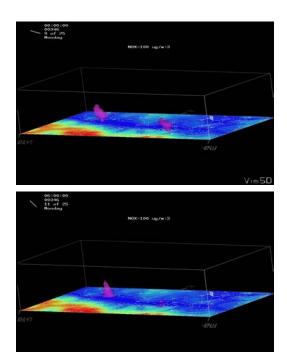
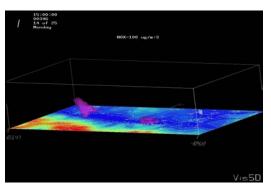


Fig. 10. RAMS/HYPACT model verification. The horizontal axis represents station number, and the vertical axis represents observed and predicted SO₂ concentration (unit: mg/m³).

5.4. Predicted Temporal and Spatial Variability of Air-Pollutants

The RAMS/HYPACT model predicts the air-pollutant concentrations. Here, we show an NOx spreading event from 07h Dec 11 to 07h Dec 12, 2000 as an illustration. Two NOx plumes (concentration > 0.1 mg/m³) occur at 07h Dec11. On Dec 11, the atmosphere has weak stratification. The plumes disperse to high altitudes. In the morning of Dec 12 (07 h), the stratification strengthens. The two plumes spread horizontally in the valley and cause high NOx concentration.





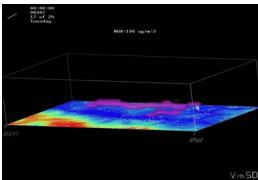


Fig. 11. Temporal variation of NOx plumes (concentration $> 0.1 \text{ mg/m}^3$): (a) 7 h, Dec 11, (b) 13h, Dec 11, (c) 22h, Dec 11, and (d) 7h, Dec 12, 2000.

6. Conclusions

We obtain two major conclusions.

- (1) Mountain-slope afforestation improves the air quality through destabilizing the atmosphere, enhancing the upward motion over the valley, and providing sinks for pollutants.
- (2) RAMS-HYPACT has capability to predict the transport of pollutants.

References

Avissar, R., 1995, Recent advances in the representation of land-atmosphere interactions in general circulation models. *Rev. Geophys., Suppl.*, 1005-1010.

Avissar, R., and R.A. Pielke, 1989, A parameterization of heterogeneous land surfaces for atmospheric numerical models and its impact on regional meteorology. Mon. Wea. Rev., 117, 2113-2136.

Chu, P.C., 1987, An icebreeze mechanism for an ice divergence-convergence criterion in the marginal ice zone. *J. Phys. Oceanogr.*, **17**, 1627-1632, 1987.

Hsu, S.A., 1973, Dynamics of the sea breeze in the atmospheric boundary layer: A case study of the free convection regime. *Mon. Weather Rev.*, **101**, 187-194.

Hu, Y., Y. Gao, J. Wang, G. Ji, Z. Shen, L. Cheng, J. Chen, and S. Li, 1994, Some achievements in scientific research during HEIFE. *Plateau Meteor.*, 13, 225-236 (in Chinese with English abstract).

Kai, K., M. Matsuda, and R. Sato, 1997, Oasis effect observed at Zhangye Oasis in the Hexi Corridor, China. *J. Meteor. Soc. Japan*, **75**, 1171-1178.

Lin, N.F., J. Tang, F.X. Han, 2001, Eco-environmental problems and effective utilization of water resources in the Kashi Plain, western Terim Basin, China. *Hydrol. J.*, **9**, 202-207.

Liu, J., and K. Kotoda, 1998, Estimation of regional evapotranspiration from arid and semi-arid surfaces, *J. Amer. Water Res. Assoc.*, **34**, 27-41.

Pielke, R.A., 2001, Influence of the spatial distribution of vegetation and soils on the prediction of cumulus convective rainfall. Rev. Geophys., 39, 151-177.

Roy, B, and R. Avissar, 2000, Scales of response of the convective boundary layer to land-surface heterogeneity. *Geophys. Res. Lett.*, **27**, 533-536.

Segal, M., and R.W. Arritt, 1992, Non-classical circulations caused by surface sensible heat-flux gradients, Bull. *Am. Meteorol. Soc.*, 73, 1593-1604.

Walko, R.L., C.J. Tremback, and M.J. Bell, 2001: HYPACT Version 1.2.0 User Guide, pp39, Mission Research Corporation, Fort Collins, CO.