

Statistical characteristics of irreversible predictability time in regional ocean models

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Abstract. Probabilistic aspects of regional ocean model predictability is analyzed using the probability density function (PDF) of the irreversible predictability time (IPT) (called τ -PDF) computed from an unconstrained ensemble of stochastic perturbations in initial conditions, winds, and open boundary conditions. Two-attractors (a chaotic attractor and a small-amplitude stable limit cycle) are found in the wind-driven circulation. Relationship between attractor’s residence time and IPT determines the τ -PDF for the short (up to several weeks) and intermediate (up to two months) predictions. The τ -PDF is usually non-Gaussian but not multi-modal for red-noise perturbations in initial conditions and perturbations in the wind and open boundary conditions. Bifurcation of τ -PDF occurs as the tolerance level varies. Generally, extremely successful predictions (corresponding to the τ -PDF’s tail toward large IPT domain) are not outliers and share the same statistics as a whole ensemble of predictions.

1 Introduction

Estimate of predictability skill of regional ocean models is important but difficult. The regional ocean dynamics is sensitive to variability of external forcing such as winds, fresh water discharge, buoyancy flux, and to multi-scale interactions between shelf and abyssal currents (Robinson et al., 1996; Lozano et al., 1996; Jiang and Malanotte-Rizzoli, 1999; Chu et al., 1999a, b, c, d, 2001; Auclair et al., 2003 among others). Bathymetry, coastlines, and physical processes such as wind bursts, fresh water discharge, tides, storm surges, shelf waves, and nonlinear processes affect coastal and abyssal currents and may cause multi-attractor circulation in marginal seas where robust circulation regimes (attractors) are more complicated than simply periodic or even quasi-periodic ones.

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Multiple attractors in atmospheric models are used by several investigators to understand the dynamics of planetary flow regimes (Charney and De Vore, 1979; Molteni, 1996). Multi-modality is found in wind-driven ocean circulation models (Schmeits and Dijkstra, 2000) such as the Kuroshio path variation south of Japan (Masuda et al., 1999). The multi-modality in the probability density function (PDF) of prediction error is caused by the spatial phase organization of the local error growth rate (Miller and Ehret, 2002). Usually, the multi-modality represents the intermittency of the predictability of atmospheric and ocean models and the formation of non-Gaussian PDF of the prediction error.

Even for prediction error with initial Gaussian PDF, the PDF of predictability time scale is often non-Gaussian with a tail stretching into large predictability time scale values (Chu and Ivanov, 2002, 2004a, b; Chu et al., 2002a, b; Denholm-Price, 2003). Rare forecasts of enhanced predictability dominate these tails. Obviously, if the occurrences of these predictions increase, the ensemble prediction will be improved.

In this study, a wind-driven circulation model based on the barotropic version of Princeton Ocean Model (POM) (Blumberg and Mellor, 1987) is implemented for a semi-enclosed rectangular basin with flat bottom to reproduce the multi-attractor regime and to study the variability of model predictability-skill with uncertain initial conditions, external forcing (wind) and open boundary conditions. A recently proposed irreversible predictability time (IPT) τ (Ivanov et al., 1994; Chu et al., 2002a, b) is taken as a quantitative measure of model predictability skill.

There are three major tasks in this study: (1) to find mechanisms for the formation of non-Gaussian PDF of IPT (called τ -PDF) with a tail stretching into large IPT domain, (2) to detect the effect of stochastic perturbations in external forcing on the model predictability, and (3) to identify the control parameter (or control parameters) for the variability or bifurcation of the model predictability skill such as the change of symmetric τ -PDF into asymmetric τ -PDF with a long tail toward large IPT domain. With fulfilling these tasks, the following fundamental questions will be answered: Do

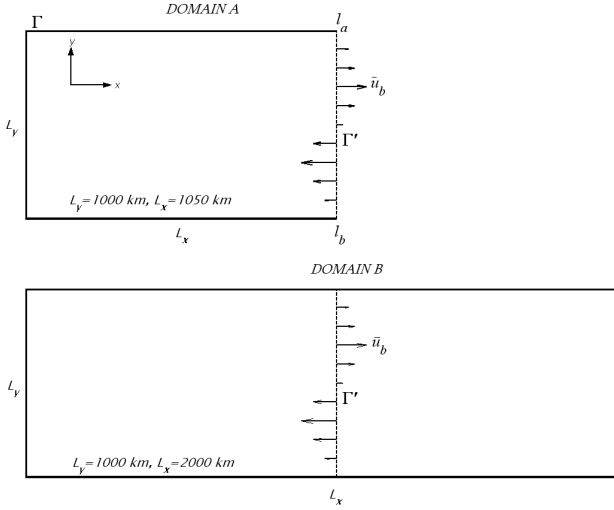


Fig. 1. Two areas for the POM integration: (a) domain-A with three rigid boundaries and one open boundary Γ' , and (b) domain-B with four rigid boundaries.

extremely successful predictions (long IPT) share the same statistical properties as predictions of short and intermediate durations? How do the stochastic perturbations in external forcing switch the attractors (noise-induced escapes)? How do they affect a multi-attractor system? What is the most important parameter to determine the τ -PDF characteristics?

The paper is organized as follows. Section 2 describes the reference solution presenting two circulation attractors: a small-amplitude stable limit cycle and a chaotic attractor. Section 3 discusses the method for generating unconstrained ensembles of perturbations added to the initial conditions, wind forcing, and open boundary conditions. Sections 4 and 5 discuss the first and second kinds of predictability. Section 6 presents the conclusions.

2 The reference solution

2.1 Numerical modeling

2.1.1 Two domains

Consider a flat bay centered at 35° N and bounded by three rigid boundaries (Domain-A in Fig. 1a). This bay expands 1000 km (1050 km) in the north-south (east-west) direction. The northern, southern, and western boundaries are rigid, and the eastern boundary is open. The Cartesian coordinate system is chosen with the origin at the southwest corner. The x - and y -axes point towards the east and north, respectively (Fig. 1). The eastern boundary of Domain-B is connected to a domain with three rigid boundaries forming a closed rectangular domain (Fig. 1b), called Domain-B (Chu et al., 1997).

The circulation in the bay is modeled with the barotropic version of POM. The vertical mixing and vertical viscosity are not included. The horizontal mixing is parameterized by

the bi-harmonic operator. The bottom stress is parameterized using the quadratic drag relation. The POM was evaluated thoroughly using the observational data (Chu et al., 2001).

2.1.2 Integration over the closed domain (Domain-B)

The POM model was integrated for Domain-B with four rigid boundaries from rest and zero surface elevation ($\hat{\xi}$),

$$(\hat{u}, \hat{v}, \hat{\xi}) = 0, \quad (1)$$

and forced by the zonal surface pseudo wind stress varying with latitude

$$\frac{W_S}{\rho_a} = -w_0 \cos\left(\frac{\pi y}{L_y}\right), \quad (2)$$

where $w_0 = 10^{-3} \text{ m}^2 \text{ s}^{-2}$, is the mean pseudo wind stress, and $\rho_a = 1.29 \text{ kg m}^{-3}$, is the air density at the ocean surface. The time step is chosen as 2 min. The horizontal resolution is 50 km.

2.1.3 Integration over the open domain (Domain-A)

We integrate the POM over Domain-B with four rigid boundaries (known boundary conditions) from the initial conditions (1) and surface forcing (2) and take the solution along the middle of Domain-B ($x = 1050 \text{ km}$)

$$(\bar{u}, \bar{\xi})_b = (\hat{u}, \hat{\xi})_b, \quad (3)$$

as the reference open boundary condition for the Domain-A integration. The velocity at $x = 1050 \text{ km}$ for the Domain-B run is nearly zonal. The solution for the Domain-B integration at day-10 are taken as the initial and open boundary conditions for the Domain-A integration,

$$(\bar{u}, \bar{v}, \bar{\xi})_{t=t_0} = (\hat{u}, \hat{v}, \hat{\xi})_{t=10 \text{ day}}. \quad (4)$$

The model for Domain-A is integrated with the wind forcing (2), open boundary condition (3), and from the initial condition (4) for another 750 days. The initial condition is an unclosed single gyre with velocities are under 0.35 m s^{-1} and sea surface elevation around 0.05 to 0.1 m. Figure 2b shows the normal velocity along the open boundary.

2.2 Two-attractor circulation

Two different attractors in the numerical solutions with the same external forcing and dissipation are detected by the temporal variability of total kinetic energy, Lyapunov dimensionality, and local Lyapunov exponents (Anishenko, 1997).

An attractor can be detected either in the state space (Berloff and McWilliams, 1999) or in the phase space. As indicated in Appendix-A, the compressible flow (u, v) with small sea surface elevation can be represented approximately by the geostrophic stream function Ψ (see Eq. A2),

$$\Psi(x, y, t) = \Psi_H(x, y) + \sum_{k=1}^K A_k(t) \Psi_k(x, y), \quad (5)$$

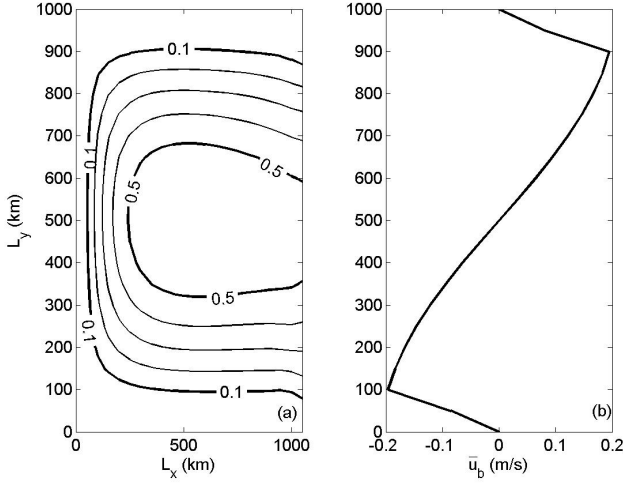


Fig. 2. Initial condition for modeling (a) and open boundary condition \bar{u}_b .

where Ψ_H is the harmonic function explained in Appendix A.

The orthonormal functions $\{\Psi_k\}$ (modes) generate a K -dimension phase space. The series (5) is truncated at $K=100$. An attractor is characterized by its basin volume and stability.

2.2.1 Basin volume

If the model dynamics is high-dimensional and dissipative, it is expected that generally the attractor basin is not a smoothly connected object, but a fractal or a riddled basin (Guckenheimer and Holmes, 1983). Accordingly to Kaneko (1998), the basin volume of an attractor is estimated as the portion of the randomly selected initial points that are attracted to it. After model integration over 100 days with 10 000 initial conditions, the attractor is identified for the model trajectory falls. Perturbed model trajectories are approximately distributed between two attractors in the proportion as 1/9.

A manifold of initial conditions $\tilde{A}_k(t_0)$ corresponding to one of the attractors with the basin volume occupied approximately 10% of phase space (the attractor is identified below as a small amplitude limit cycle) can be specified through smoothing (filtrating) initial conditions accordingly to

$$\tilde{A}_k(t_0) = \frac{\lambda_k^2}{\lambda_k^2 + \chi_k^2} A_k(t_0), \quad (6)$$

where the function χ_k is defined by

$$\chi_k = \chi_0 \max |A_k(t_0)| / A_k(t_0),$$

and χ_0 is the filter parameter; λ_k is the eigenvalue corresponding to the basis function $\Psi_k(x, y)$. Chu et al. (2003) demonstrate that the filter (6) removes high-order modes and low-order small-amplitude modes.

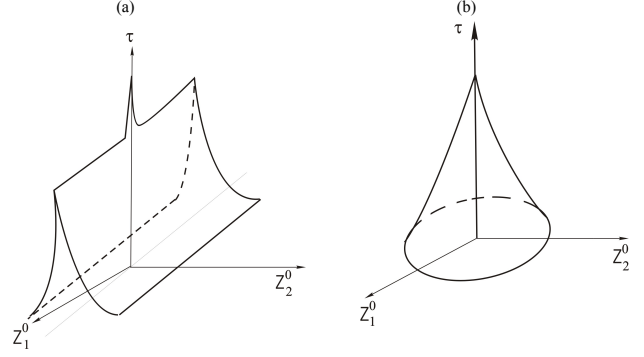


Fig. 3. Typical IPT landscape projected onto error phase space $\{Z_1^0, Z_2^0\}$: (a) chaotic attractor and (b) limit cycle.

2.2.2 Measure of attractor stability

To estimate attractor stability in forecast metrics

$$J(t) = \frac{\|\mathbf{Z}^T \mathbf{Z}\|^2}{\|\bar{\Psi}\|^2}, \quad (7)$$

$$\mathbf{Z}(\mathbf{x}, t) = \Psi(\mathbf{x}, t) - \bar{\Psi}(\mathbf{x}, t), \quad \mathbf{Z}(\mathbf{x}, t_0) = \mathbf{Z}_0, \quad (8)$$

where (\mathbf{x}, t) denotes the spatial and temporal coordinates; $\bar{\Psi}(\mathbf{x}, t)$ and $\Psi(\mathbf{x}, t)$ are the reference solution and an individual prediction, respectively; the superscript ‘T’ denotes the transpose operator, $\|\cdot\|$ is the Euclidian norm, we use the irreversible predictability time determined as

$$\tau = \min \left(t \mid J(t) > \varepsilon^2 \right), \quad (9)$$

where $\varepsilon = \varepsilon_0 / \|\Psi\|^2$, ε_0 is tolerance level (accepted prediction accuracy).

IPT differs from the e-folding time (τ_e) or doubling time if prediction error is oscillatory or stochastic. For example, the mean IPT over an ensemble of initial conditions is defined by

$$\langle \tau \rangle = \left\langle \min \left(t \mid J(t) > \varepsilon^2 \right) \right\rangle. \quad (10)$$

However, the e-folding time is determined by

$$\langle \tau_e \rangle = \min \left(t \mid \langle J(t) \rangle > \varepsilon^2 \right). \quad (11)$$

with the corresponding ε^2 .

Clearly, unique correlations exist among the landscape of IPT (dependence of IPT on magnitude and orientation of the initial error), the local Lyapunov exponents and the attractor types. In particular, Chu and Ivanov (2004a, b) pointed out that for a given IPT landscape, the signs of the local Lyapunov exponents can be determined.

For a chaotic attractor with at least one positive exponent and one negative exponent, the IPT landscape has a crest-like structure (Fig. 3a). If the initial error is oriented perpendicular to the crest, the prediction error grows exponentially. If the initial error is oriented along the crest, the most unstable

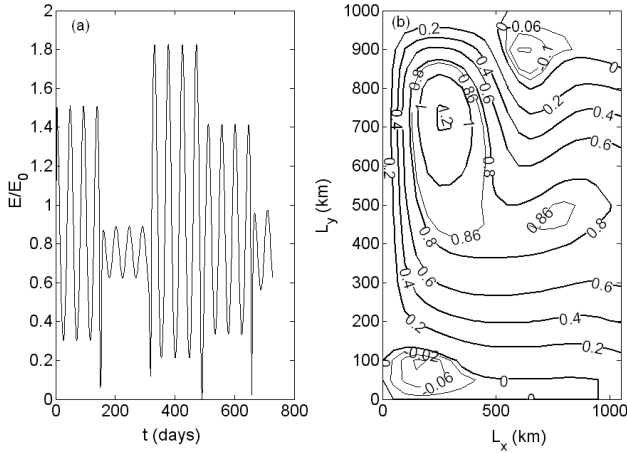


Fig. 4. The chaotic attractor: **(a)** mean kinetic energy normalized by the value of maximum mean kinetic energy of limit cycle E_0 ; **(b)** snapshot of stream function corresponding to a maximum energy state of attractor.

Lyapunov vector decays (stagnates). A hill-like landscape corresponds to a stable limit cycle (Fig. 3b). Here, there is no specified orientation in the phase space.

2.3 Attractor patterns

2.3.1 Chaotic attractor

After integrating the model from the original (taken as accurate) initial condition (Fig. 2b), the attractor is characterized by a-periodic behavior of kinetic energy (Fig. 4a). The streamfunction with the maximum energy (Fig. 4b) shows a strong cyclonic basin-scale gyre with size of about 350 km (left upper corner) and speed of about 0.9–1.0 m s⁻¹. Several smaller size (secondary) eddies occur near the rigid boundary and in the center of the basin. The highest surface elevation reaches 1 m.

The nonlinear regime is identified in the basin by two non-dimensional Rossby numbers (Pedlosky, 1987),

$$Ro_1 = \max\left(\frac{U}{fL}\right) \sim 0.1, \quad Ro_2 = \max\left(\frac{1}{fT}\right) \sim 0.12 \quad (12)$$

where $f \sim 10^{-4} \text{ s}^{-1}$ is the Coriolis parameter; $U \sim 1 \text{ m s}^{-1}$ is the characteristic velocity in the basin; $L \sim 100 \text{ km}$ and $T \sim 1 \text{ day}$ are the characteristic spatial and temporal scales of the flow, respectively.

Temporal variation of the kinetic energy averaged over the whole basin (Fig. 4a) shows a-periodic oscillations and transitions between meta-stable states. These oscillations are caused by interactions between the basin-scale cyclonic gyre and the secondary eddies which are generated by barotropic instability of the gyre when the circulation velocity exceeds 0.7–0.8 m s⁻¹. The instability produces the secondary eddies disposing along the rigid boundary and in the center of the basin and periodically dissipating due to viscosity damping. Note that for zero horizontal viscosity the chaotic attractor

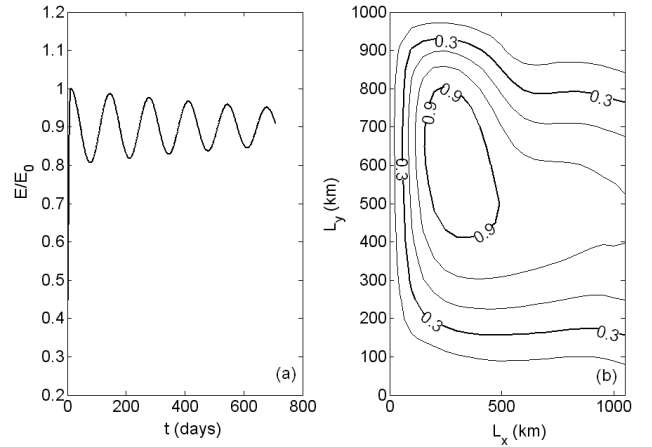


Fig. 5. The small-amplitude limit cycle: **(a)** mean kinetic energy normalized by its maximum value E_0 , **(b)** snapshot of stream function averaged over time.

does not develop. Instead an unstable limit cycle occurs. This case will be discussed in Chu and Ivanov (2004a, b).

The analysis of IPT landscape (Chu and Ivanov, 2004a, b) demonstrates the existence of at least one positive and one negative Lyapunov exponent. It identifies the observed dynamical regime as a chaotic attractor.

2.3.2 Stable limit cycle

If the nonlinear filtration procedure (4) is applied to the initial condition (4), an oscillated stable cyclonic gyre quickly develops for the semi-enclosed basin. Figure 5a shows the temporal variation of the kinetic energy averaged over the whole basin. The oscillation period is about 150 days. The nondimensional amplitude of the oscillation is around 0.1. Figure 5b shows temporally averaged stream function.

Such a circulation pattern is represented as a stable limit cycle (small amplitude) in the phase space. This limit cycle is identified using the method proposed by Berloff and Meacham (1997).

3 Unconstrained ensemble of perturbations

Since the barotropic model is computationally efficient, the stochastic stability of the reference solution to unconstrained perturbations of the initial condition, open boundary condition, and wind forcing is studied using the Monte-Carlo method with 1000 ensemble samples.

Atmospheric ensemble modeling shows that for small samples the initial error can not cover the phase space homogeneously. It leads to an unrealistically-weak growth of prediction error variance. Therefore, for short period prediction it requires the best possible knowledge of the initial state in the sub-region of the phase space where the prediction errors are most likely to grow. In different forecast centers, such as the European Centre for Medium-Range Weather Forecasts

and the National Centers for Environmental Prediction, the directions corresponding to the maximum error growth in the phase space are identified by the singular vectors (Buizza et al., 1999) and the breed vectors (Toth and Kalnay, 1997).

This study utilizes a simple Monte-Carlo technique for the three reasons. First, little difference in the statistics is found among ensembles of 1000, 5000, 10 000 and 20 000 samples. The numerical experiments with 1000 ensemble sample are reliable. Second, a Latin hypercube sampling procedure (LHST, 2001) is used to increase the degree of uniform coverage of the phase space by initial errors. Third, a non-parametrical technique using the Epanechnikov kernel is used to reconstruct PDF from histogram. Details of such a reconstruction can be found in Good (1996).

3.1 Perturbations to initial condition

Sensitivity of the reference solution to Gaussian-type perturbations is commonly studied in the atmospheric and oceanic applications (Robinson et al., 1996). In this study, 2D stochastic perturbations with the correlation function

$$G(\mathbf{r} - \mathbf{r}') = I^2 \exp \left[-\frac{(\mathbf{r} - \mathbf{r}')^2}{R^2} \right] \quad (13)$$

are used. Here, $\mathbf{r} = x\mathbf{i} + y\mathbf{j}$, is the displacement vector; (\mathbf{i} , \mathbf{j}) are unit vectors along x and y axes; (R , I) are the correlation radius and intensity of the noise. The technical details of generating these noises can be found in Sabel'feld (1991).

3.2 Perturbations to wind forcing

Surface winds are perturbed by

$$\mathbf{u}' \equiv [u'(x, y, t), v'(x, y, t)] = [\mu_1(t), \mu_2(t)]g(x, y), \quad (14)$$

where μ_1 and μ_2 are independent white noise processes with zero mean and variance of σ^2 . In numerical integration, the noises (μ_1, μ_2) are introduced every hour. The spatial structure function of errors, g , is parameterized by,

$$g(x, y) = \alpha \left[\pi \lambda_x \lambda_y \operatorname{erf} \left(\frac{L_x}{2\lambda_x} \right) \operatorname{erf} \left(\frac{L_y}{2\lambda_y} \right) \right]^{1/2} \exp \left(-\frac{(x - x_0)^2}{\lambda_x^2} - \frac{(y - y_0)^2}{\lambda_y^2} \right), \quad (15)$$

which describes the impact of the localized atmospheric eddy activity near (x_0, y_0) on the surface wind perturbations (Sura et al., 2001). Here, erf is the error function; α is a scaling parameter; (λ_x , λ_y) are the de-correlation scales. Noise in surface winds with $\sigma^2 = 28 \text{ m}^2 \text{ s}^{-2}$ corresponds to observed typical atmosphere conditions in the North Atlantic region (Wright, 1988).

3.3 Perturbations to open boundary condition

The normal velocity of the reference solution along the open boundary (u_b) is perturbed by

$$u'_b = B(t) \sin(\pi y / L_y), \quad (16)$$

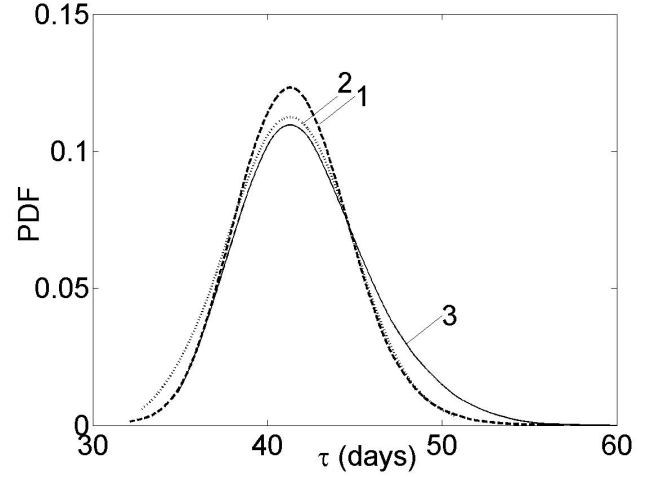


Fig. 6. τ -PDF for the predictability of the first kind with white initial noise (curve-2), and red noise with the correlation radius of 14 km (curve-1) and 70 km (curve-3). Here, $I^2 = 0.01 \text{ m}^2/\text{s}^2$ and $\varepsilon_0^2 = 0.5$.

where $B(t)$ is white noise with zero mean and variance of δ^2 . The velocity perturbation (u'_b) affects inflow and outflow across the open boundary.

4 Predictability of the first kind

Keeping wind forcing (2) and open boundary condition (4) accurate, and perturbing initial condition (3) with white or red noise represented by two parameters (variance of initial error I^2 and correlation radius R), the predictability of the first kind (Lorenz, 1982; Chu, 1999) can be investigated with various combinations of (I^2 , R , ε_0^2). In the phase space, perturbed model trajectories start from the basin of chaotic attractor, i.e. from the original initial condition (3). The τ -PDF is sensitive to the variation of ε_0^2 .

For a given tolerance level ε_0^2 (0.5), the τ -PDF is close to Gaussian shape (skewness of 0.08) for the white noise ($R=0$, curve-1 in Fig. 6). Clearly, any model trajectory starting from the attractor basin does not leave the attractor and the IPT statistics is only determined by the stability of this chaotic attractor.

When the correlation radius R increases, the volume in the phase space covered by the initial error increases (Chu and Ivanov, 2004a), and the two-attractor structure of the reference circulation occurs. For a given error intensity, $I^2 = 0.01 \text{ m}^2 \text{ s}^{-2}$, the τ -PDF is near Gaussian (curve-1 in Fig. 6, skewness near 0) for $R=0$ km, non-Gaussian (curves-2 in Fig. 6, skewness of 0.58) for $R=14$ km, and non-Gaussian with a tail stretching to large IPT domain (curves-3 in Fig. 6, skewness of 0.80) for $R=70$ km.

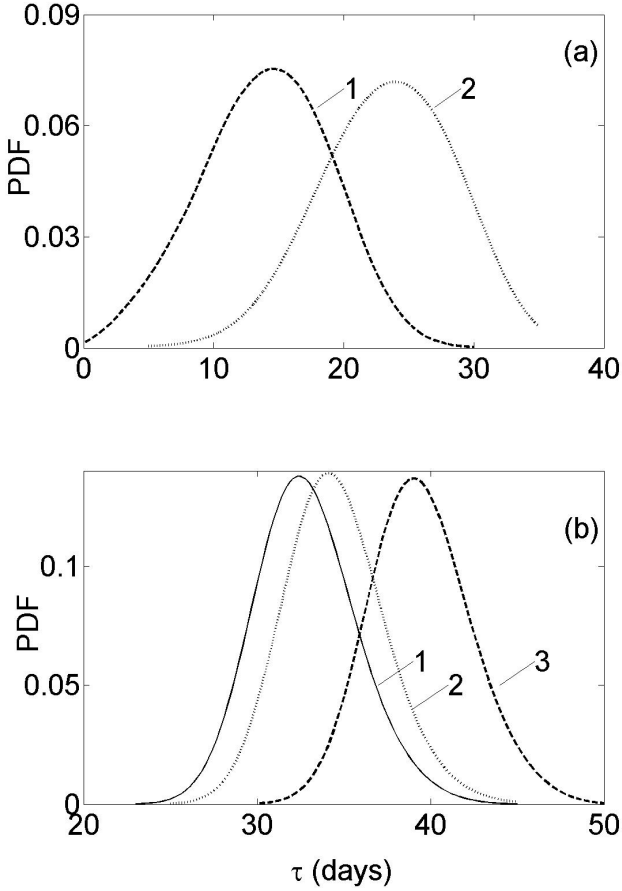


Fig. 7. τ -PDF for the predictability of the second kind (wind forcing) with a given wind error intensity ($\sigma^2=42 \text{ m}^2 \text{ s}^{-2}$) and various tolerance levels (ε_0^2): **(a)** small tolerance levels: 0.005 (curve-1), 0.01 (curve-2), and **(b)** intermediate and large tolerance levels: 0.05 (curve-1), 0.1 (curve-2) and 0.5 (curve-3).

5 Predictability of the second kind

Keep the initial condition (3) accurate, and perturb the wind forcing (2) and open boundary condition (4) with white noises. The noise the intensity is σ^2 for the wind forcing and δ^2 for the open boundary condition. The second kind of predictability (Lorenz, 1982) can be investigated with various combinations of (σ^2 , ε_0^2) and (δ^2 , ε_0^2). Since the predictability of high-resolution regional models depends crucially on the wind forcing (Robinson et al., 1996; Chu et al., 1999a, b, c; Burillo et al., 2002 and others) and the specified normal velocity at the open boundaries, and since errors in the open boundary condition cause a rapid decrease of model predictability skill (Chu, 1999; Jiang and Malanotte-Rizzoli, 1999), investigation of the second kind of predictability is of great importance for regional ocean modeling.

5.1 Residence time and τ -PDF

In contrast to the predictability of the first kind, analysis of the second kind of predictability is more difficult because of

the multi-attractor system. “Noise induced escape” occurs when a model trajectory switches over attractors (Anishenko, 1997). Therefore, in addition to IPT, the residence time (τ_{res}) should be used with the definition of a maximum time period for a perturbed trajectory keeping in the same attractor.

In general, the residence time depends on the basin volume of the attractor, initial position of model trajectory at the attractor, dynamical attractor stability and noise intensity (σ^2 or δ^2) (Kaneko, 1997). Since there is no common method to estimate τ_{res} on the attractor for a high-dimensional dissipative dynamical system (Kaneko, 1997), a practical method is proposed.

For a small ε_0^2 , it is expected that

$$\tau < \tau_{res},$$

which shows that the model loses the predictability skill before the model trajectory leaves the occupied attractor. The τ -PDF is determined by the stability of a single attractor and should be near Gaussian type. For a sufficiently large ε_0^2 , it is expected that

$$\tau > \tau_{res},$$

which shows that the model keeps the predictability skill after the model trajectory leaves the occupying attractor. This leads to the change of the τ -PDF structure. Therefore, switch among attractors can be determined through analyzing the change of τ -PDF structure with increasing ε_0^2 .

The proposed methodology, of course, can not determine the residence time τ_{res} and the number for a trajectory switching among attractors. However, it determines the change of the τ -PDF structure that links to the duration of the perturbed trajectory along different attractors. Besides, it uniquely identifies the trajectory switching from the chaotic attractor to the limit cycles when the stochastic perturbation is added to external forcing.

5.2 Perturbations to wind forcing

The two parameters σ^2 and ε_0^2 determine the τ -PDF structure. For small tolerance level ($\varepsilon_0^2 < 0.01$) and strong wind perturbations ($\sigma^2 > 40 \text{ m}^2 \text{ s}^{-2}$), the τ -PDF is nearly symmetric with large variance of IPT. The kurtosis and skewness are approximately 3.1 and 0.1 (Fig. 7a).

With increase of ε_0^2 (> 0.075), the τ -PDF becomes asymmetric with smaller IPT variance and a long tail stretching to large IPT domain (Fig. 7b). The skewness is 0.60 for curve-1 and 0.75 for curve-3. This shows that the tolerance level ε_0^2 is a controlling parameter for model predictability skill. The τ -PDF is symmetric for small ε_0^2 and asymmetric for large ε_0^2 . With a relatively large tolerance level ($\varepsilon_0^2=0.1$), the asymmetry of the τ -PDF structure increases (longer tail toward large IPT domain) with the increase of noise intensity σ^2 (Fig. 8). The extra-successful predictions (Chu et al., 2002c) corresponding to the PDF tail (toward the large IPT domain) are not outliers since they share the same statistics with other predictions with short and intermediate IPTs.

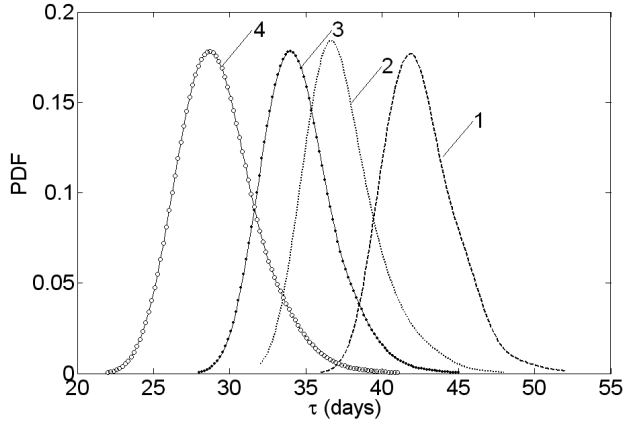


Fig. 8. τ -PDF for the predictability of the second kind (wind forcing) with a given tolerance level $\varepsilon_0^2=0.1$ with various wind error intensity: $\sigma^2=14\text{ m}^2\text{ s}^{-2}$ (curve-1), $28\text{ m}^2\text{ s}^{-2}$ (curve-2), $42\text{ m}^2\text{ s}^{-2}$ (curve-3), and $56\text{ m}^2\text{ s}^{-2}$ (curve-4).

Another important result is that the multi-attractor structure of the modeled circulation does not necessary lead to the multi-modal τ -PDF since the perturbed trajectory can switch among different attractors. This is different from a recent result reported by Miller and Ehret (2002) that if the initial error co-variances are large enough such that a given perturbed trajectory has high possibility to start from different basin attractions. The initial Gaussian τ -PDF will evolve into a bimodal τ -PDF.

Existence of a tail stretching into large IPT domain is the most common feature of τ -PDF for short and intermediate forecasts (Fig. 8). Increase of the noise intensity σ^2 enhances the asymmetry of the τ -PDF. For example, the skewness is 0.38 for $\sigma^2=14\text{ m}^2\text{ s}^{-2}$ (curve-1) and 0.96 for $\sigma^2=56\text{ m}^2\text{ s}^{-2}$ (curve-4). The extremely successful predictions (Chu et al., 2002b) corresponding to the τ -PDF tails are not outliers since they share the same statistics with the usual predictions. This will be pointed out in Chu et al. (2004b) that the Weibull distribution is the best fit for the τ -PDFs illustrated in Fig. 8.

The τ -PDF is nearly symmetric for small ε_0^2 and asymmetric for large ε_0^2 . Loss of the τ -PDF symmetry occurs as $\varepsilon_0^2 > 0.075$. This indicates that variation of the tolerance level ε_0^2 may lead to a drastic change of statistics for measuring the model predictability skill. A set of structural change of the τ -PDF with the increase of ε_0^2 may be considered as the one-step bifurcation sequence:

$$\tau\text{-PDF symmetry} \rightarrow \tau\text{-PDF asymmetry.} \quad (17)$$

5.3 Perturbations to open boundary condition

For small tolerance level ($\varepsilon_0^2 \leq 0.01$), the τ -PDF is already asymmetric with a maximum value at small IPT domain (about 5 days, i.e., the most probable IPT) and a tail stretching to large IPT domain up to 40–50 days (curves-1 and curve-2 in Fig. 9). The skewness is 3.8 for $\varepsilon_0^2=0.005$ (curve-1) and 3.9 for $\varepsilon_0^2=0.01$ (curve-2). The estimated mean res-

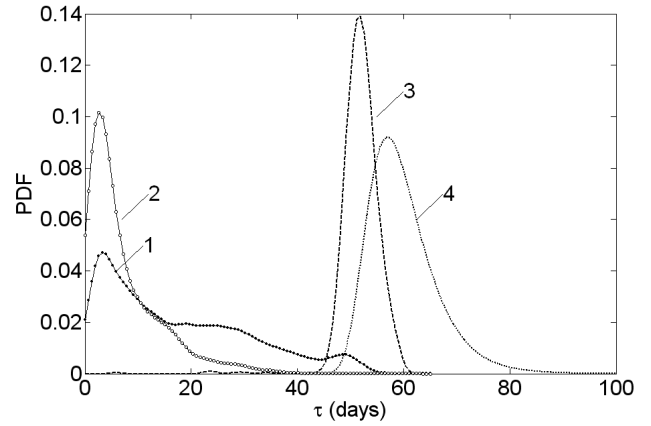


Fig. 9. τ -PDF for the second kind of predictability (open boundary) with a given open boundary error intensity ($\delta^2=0.2\text{ m}^2\text{ s}^{-2}$) and various tolerance levels (ε_0^2): 0.005 (curve-1), 0.01 (curve-2), 0.1 (curve-3) and 0.5 (curve-4).

idence time is about 10 days, the perturbed model trajectories departure quickly from the chaotic attractor to the small amplitude limit cycle. With increase of ε_0^2 (>0.025), the IPT variance and the degree of τ -PDF asymmetry reduces (curves-3 and curve-4 in Fig. 9). The skewness reduces to 0.7–0.8. This is caused by the perturbed trajectories having time to return to the chaotic attractor from the limit cycle.

For some values of ($\varepsilon_0^2, \delta^2$), there exist “outliers” – predictions with statistical properties differing considerably from those of the prediction ensemble. For example, when $\varepsilon_0^2=0.1$, $\delta^2=0.3\text{ m}^2\text{ s}^{-2}$, the τ -PDF shows a complex structure with two distinct features for $\tau < 40$ day and $\tau > 40$ day (Fig. 10a). The predictions for $\tau < 40$ day may be identified as outliers. Therefore the τ -PDF in Fig. 10a can be represented as the combination of two τ -PDFs: (a) uniform distribution (shown in Fig. 10b) for $12\text{ h} \leq \tau \leq 39$ days and (b) near Gaussian distribution with the mean and variance equal to 50 day and 64 day^2 , respectively. A set of structural change of the τ -PDF with the increase of ε_0^2 may be considered as the two-step bifurcation sequence:

$$\begin{aligned} &\text{asymmetric } \tau\text{-PDF with large variance} \rightarrow \\ &\rightarrow \text{symmetric } \tau\text{-PDF} + \text{outliers} \rightarrow \\ &\rightarrow \text{asymmetric } \tau\text{-PDF with small variance.} \end{aligned} \quad (18)$$

Existence of bifurcation sequences (17) and (18) should be checked in atmospheric and oceanic models.

6 Conclusions

(1) Several predictability regimes in two-attractor circulations (the chaotic attractor and small-amplitude limit cycle) reproduced by a numerical model have been identified in a semi-closed basin with flat bottom. The τ -PDF is non-Gaussian for intermediate and large amplitude prediction errors, and near Gaussian for small amplitude prediction errors.

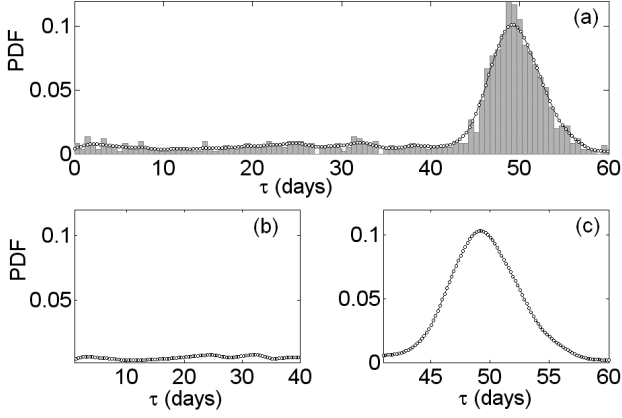


Fig. 10. “Outliers” – forecasts with $\delta^2=0.3 \text{ m}^2 \text{ s}^{-2}$, $\varepsilon_0^2=0.1$: (a) τ -histogram computed for the perturbed open boundary, (b) best fit to the homogeneous distribution for $\tau < 40$ day, and (c) best fit to Gaussian distribution for $\tau > 40$ day.

(2) For the predictability of the first kind, the error evolution depends strongly on the correlation radius of the initial error and its variance. If the initial co-variances and correlation radius are large enough so that there is large possibility that a given perturbed model trajectory does not start from the same attractor basin as the mean ensemble, the τ -PDF is asymmetric (non-Gaussian) with a tail stretching into large IPT domain.

(3) The stochastic perturbations added to the wind forcing and open boundary condition lead to switch of the perturbed model trajectories over attractors (i.e., the noise-induced escape). It is found that the perturbed model trajectories switch from the chaotic attractor to the limit cycles and switch back. Relationship between the residence time at the attractor and IPT determines the shape of τ -PDF. For $\tau < \tau_{res}$ perturbed model trajectories do not have time to switch from the chaotic attractor to the limit cycle. This causes the τ -PDF near Gaussian. For $\tau \sim \tau_{res}$ the perturbed model trajectories can be either in the initial chaotic attractor or in the limit cycle. This causes the bimodal (not often) or asymmetric non-Gaussian τ -PDF. For $\tau > \tau_{res}$, the perturbed trajectories have sufficient time to switch to the limit cycle, and cause asymmetric τ -PDF.

(4) The perturbed model trajectories attracted to the limit cycle share the same statistical properties as the trajectories still along the initial chaotic attractor. In general, the numerical computations do not indicate the outlier feature of predictions except for some forecasts realized for specific combinations of the tolerance level and variance of perturbations added to the open boundary condition.

(5) Two different bifurcation sequences in changes of τ -PDF are found with increase of ε_0^2 . For uncertain wind forcing, one-step bifurcation is found with symmetric τ -PDF changing to asymmetric τ -PDF. For uncertain open boundary conditions, a more complicated two-step bifurcation is found in the model predictability skill.

Appendix A

Spectral decomposition

A two-dimensional compressible circulation (u, v) can be represented by

$$u = -\frac{\partial \Psi}{\partial y} + \frac{\partial \Phi}{\partial x}, \quad v = \frac{\partial \Psi}{\partial x} + \frac{\partial \Phi}{\partial y}, \quad (\text{A1})$$

where Ψ and Φ are the geostrophic streamfunction and velocity potential, respectively. Let (Γ, Γ') be the rigid and open boundaries. The geostrophic streamfunction Ψ is decomposed into (Eremeev et al., 1992; Chu et al., 2003)

$$\Psi(x, y, t) = \Psi_H(x, y, t) + \sum_{k=1}^{\infty} A_k(t) \Psi_k(x, y), \quad (\text{A2})$$

where $\{\Psi_k\}$ are the eigen-functions of the spectral problem ,

$$\Delta \Psi_k(x, y) = -\lambda_k \Psi_k(x, y), \quad \Psi_k|_{\Gamma \cup \Gamma'} = 0, \quad (\text{A3})$$

with λ_k the eigen-values of the horizontal Laplacian (Δ). The component Ψ_H is the harmonic function satisfying the boundary conditions

$$\Psi_H|_{\Gamma} = 0, \quad \Psi_H|_{\Gamma'} = \Psi_b, \quad (\text{A4})$$

where Ψ_b is computed from knowledge of the normal velocity at open boundary.

The velocity potential Φ is decomposed into

$$\Phi = \sum_{m=1}^{\infty} B_m(t) \Phi_m(x, y), \quad (\text{A5})$$

where $\{\Phi_m\}$ are the eigen-functions of the spectral problem

$$\begin{aligned} \Delta \Phi_m(x, y) &= -\mu_m \Phi_m(x, y), \\ \Phi_m|_{\Gamma'} &= 0, \\ \mathbf{n} \cdot \nabla \Phi_m|_{\Gamma} &= 0, \end{aligned} \quad (\text{A6})$$

where \mathbf{n} is the unit vector normal to Γ' ; $\{\mu_m\}$ are the eigen-values.

The basis functions $\{\Psi_k, \Phi_m\}$ generate the phase space which can be used for the spectral analysis of ocean circulation in a basin with real coastlines. Since the surface elevation is small comparing to the water depth H ,

$$\zeta/H \ll 1$$

the quasi-geostrophic approximation is reasonable for the flow reproduced in the present study (Pedlosky, 1987). Thus, in the main text, only the geostrophic streamfunction Ψ is considered.

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