Adaptation of Global Volume Transport for Coastal Ocean Models

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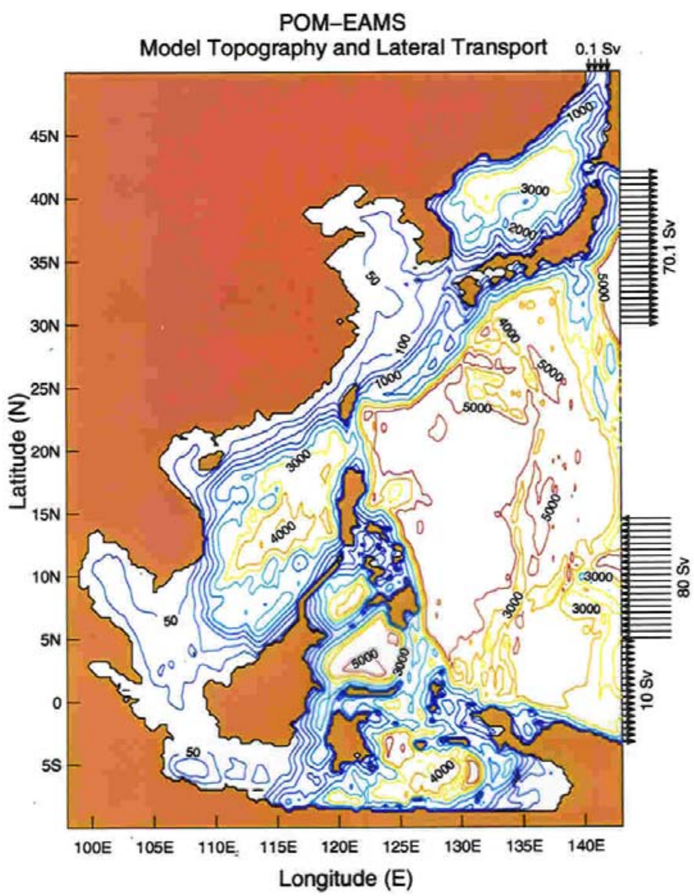
One Difficulty in Coastal Ocean Modeling:
Unknown Volume Transport at Open Boundaries

- Subjective



- Can we determine the volume transport objectively?

A possible approach

Diagnostic Method

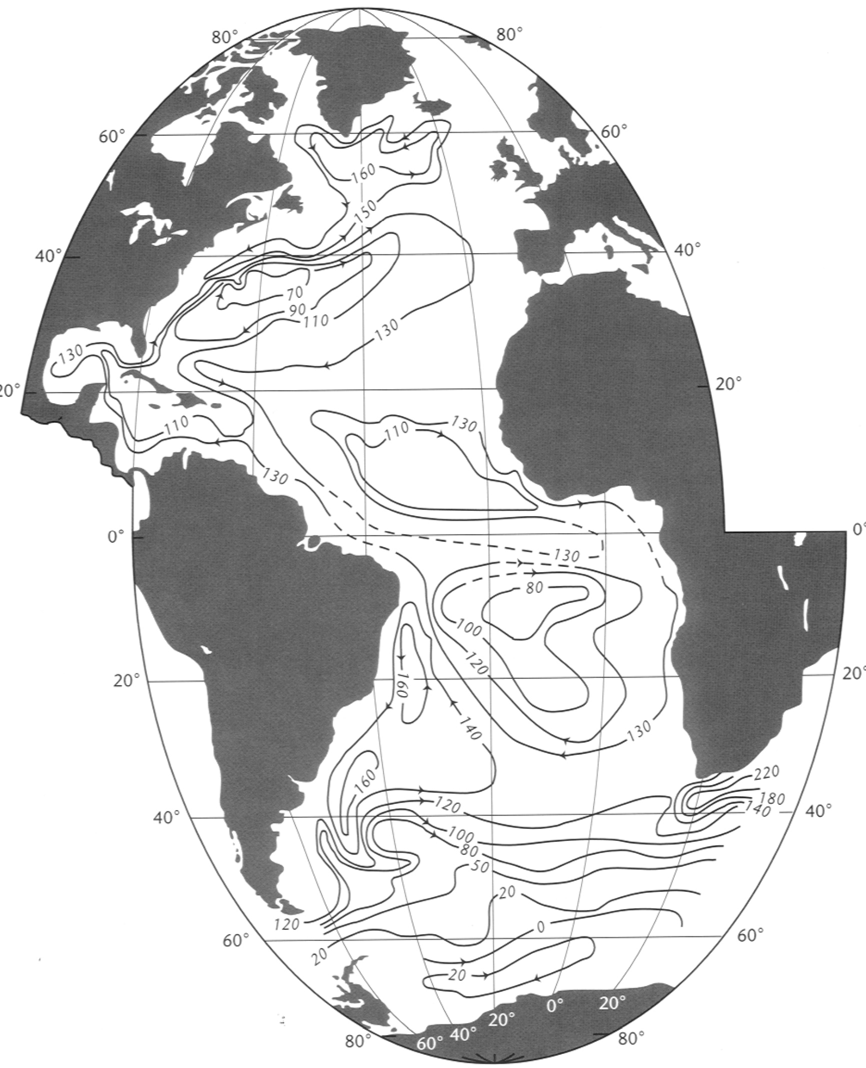
Existing Diagnostic Methods for Determination of Volume Transport

- From (T, S) Data – Total Geostrophic Flow (Reid 1985, 1994, …)

- Top 500 m Transport From (T, S) and Wind Data – Sverdrup Model (Godfrey 1989)

Reid (1994)

- No wind driven component



Three Components of the Ekman-Munk Model

- Ekman Model for Extra-Equatorial Regions
- Munk Model for the Equatorial Region
- Stokes Theorem for Determining Ψ for Islands

Steady-State Large-Scale Dynamics (Mid-Latitudes)

$$-f(v - v_g) = A_z \frac{\partial^2 u}{\partial z^2} + A_h \nabla^2 u$$

$$f(u - u_g) = A_z \frac{\partial^2 v}{\partial z^2} + A_h \nabla^2 v$$

$$\frac{\partial p}{\partial z} = -\rho g$$

$$u_g = -\frac{1}{f\rho_0}\frac{\partial p}{\partial y}, \quad v_g = -\frac{1}{f\rho_0}\frac{\partial p}{\partial x}$$

$$\nabla^2 \equiv \frac{\partial^2}{\partial x^2} + \frac{\partial^2}{\partial y^2}$$

(1) Ekman Number (Mid-Latitude)

$$E = \frac{O(|A_h \nabla^2 \mathbf{V}|)}{O(|f\mathbf{V}|)} = \frac{A_h}{|f|\, L^2}$$

- $L \sim 2 \times 10^5 m$

$$A_h = 5 \times 10^5 \ \mathrm{m^2 s^{-1}}$$

$$E \simeq \frac{5 \times 10^5 \ \mathrm{m^2 s^{-1}}}{(10^{-4}\mathrm{s^{-1}}) \times (2 \times 10^5 \mathrm{m})^2} = 0.125$$

- Horizontal diffusion can be neglected

Ekman Model for Mid-Latitudes

$$-f(v - v_g) = A_z \frac{\partial^2 u}{\partial z^2}$$

$$f(u - u_g) = A_z \frac{\partial^2 v}{\partial z^2}$$

Vertically Integrated Ekman Model

$$-f(V - V_g) = A_z \frac{\partial u}{\partial z}\Big|_{z=\eta} - A_z \frac{\partial u}{\partial z}\Big|_{z=-H}$$

$$f(U - U_g) = A_z \frac{\partial v}{\partial z}\Big|_{z=\eta} - A_z \frac{\partial v}{\partial z}\Big|_{z=-H}$$

Vertically Integrated Velocity

$$U = \widehat{U}_g + U_r + \frac{\tau_y}{f\rho_0}, \quad V = \widehat{V}_g + V_r - \frac{\tau_x}{f\rho_0}$$

Vertically Integrated Velocity

$$(\widehat{U}_g, \widehat{V}_g) = \frac{g}{f\rho_0}\left(\int_{-H}^{\eta}\int_{-H}^{z}\frac{\partial\rho}{\partial y}dz'dz, -\int_{-H}^{\eta}\int_{-H}^{z}\frac{\partial\rho}{\partial x}dz'dz\right).$$

$$U_r = Hu_{-H} - \frac{C_D}{f}\sqrt{u^2_{-H} + v^2_{-H}}u_{-H},$$

$$V_r = Hv_{-H} + \frac{C_D}{f}\sqrt{u^2_{-H} + v^2_{-H}}v_{-H}.$$

- CD ~ Bottom Drag Coefficient (0.0025)

Volume Transport Stream Function

$$U = -\frac{\partial \Psi}{\partial y}, \quad V = \frac{\partial \Psi}{\partial x}$$

Poisson-Ψ Equation

$$\nabla^2 \Psi = \Pi, \quad \Pi = \frac{\partial V}{\partial x} - \frac{\partial U}{\partial y}$$

$$\Pi = \Pi_1 + \Pi_2 + \Pi_3$$

Forcing Terms in The Poisson-Ψ Equation

- T, S

$$\Pi_1 = (\frac{\partial \widehat{V}_g}{\partial x} - \frac{\partial \widehat{U}_g}{\partial y})$$

- T, S (P-vector Inverse Method)

$$\Pi_2 = (\frac{\partial V_r}{\partial x} - \frac{\partial U_r}{\partial y})$$

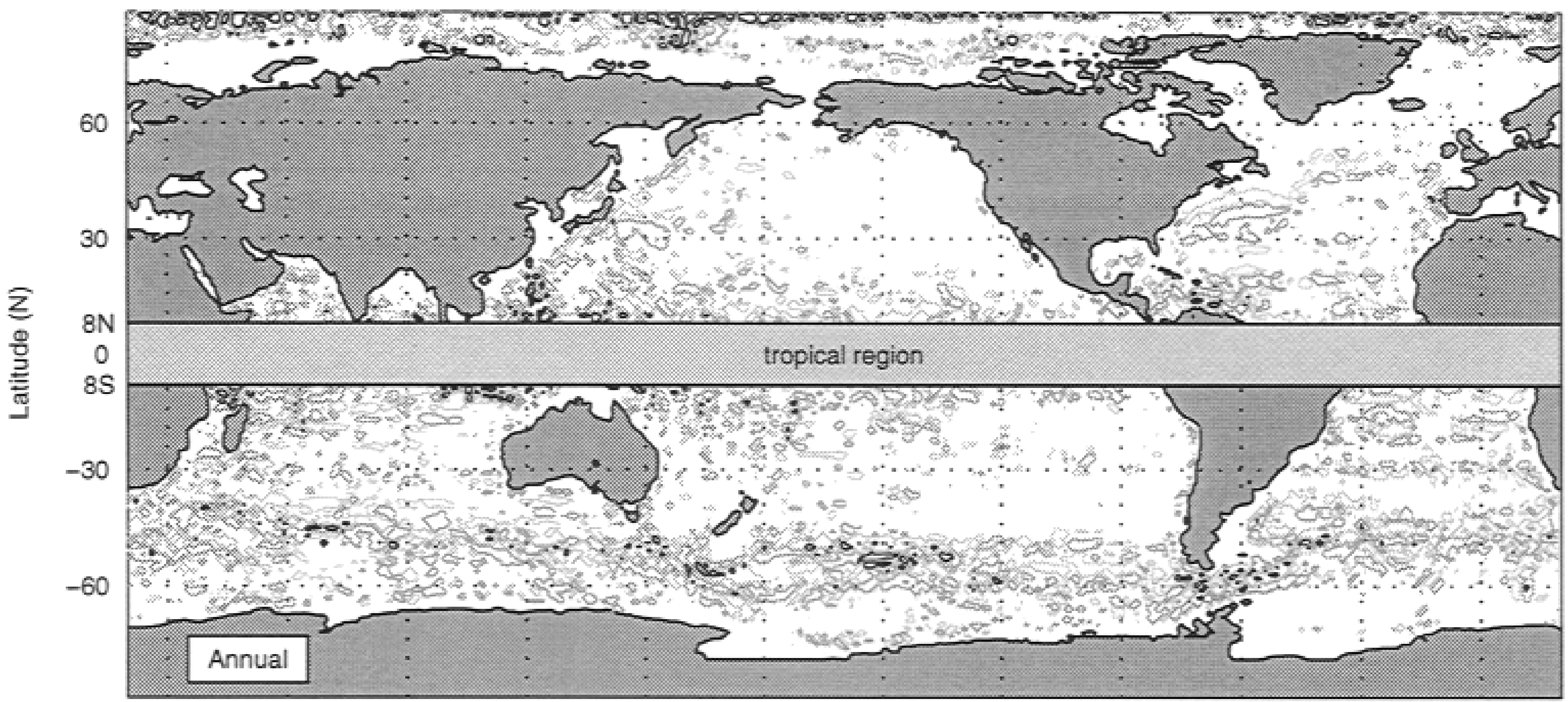
- Wind Stress

$$\Pi_3 = -[\frac{\partial}{\partial x}(\frac{\tau_x}{f\rho_0}) + \frac{\partial}{\partial y}(\frac{\tau_y}{f\rho_0})].$$

Datasets

- Topography: DBDB5, 5’ resolution
- Annual and Monthly Mean T, S: NODC (Levitus et al. 1998)
- Annual and Monthly Mean Wind Stress: NCEP

Annual Mean Π Values



Ekman Number (Low-Latitudes) 8° S – 8° N

- Horizontal diffusion cannot be neglected

$$E \geq \frac{5 \times 10^5 \ \mathrm{m^2 s^{-1}}}{(0.2 \times 10^{-4} \mathrm{s^{-1}}) \times (2 \times 10^5 \mathrm{m})^2} = 0.5$$

(2) Munk Model (Equatorial Region)

- Vertically Integrated Vorticity Equation

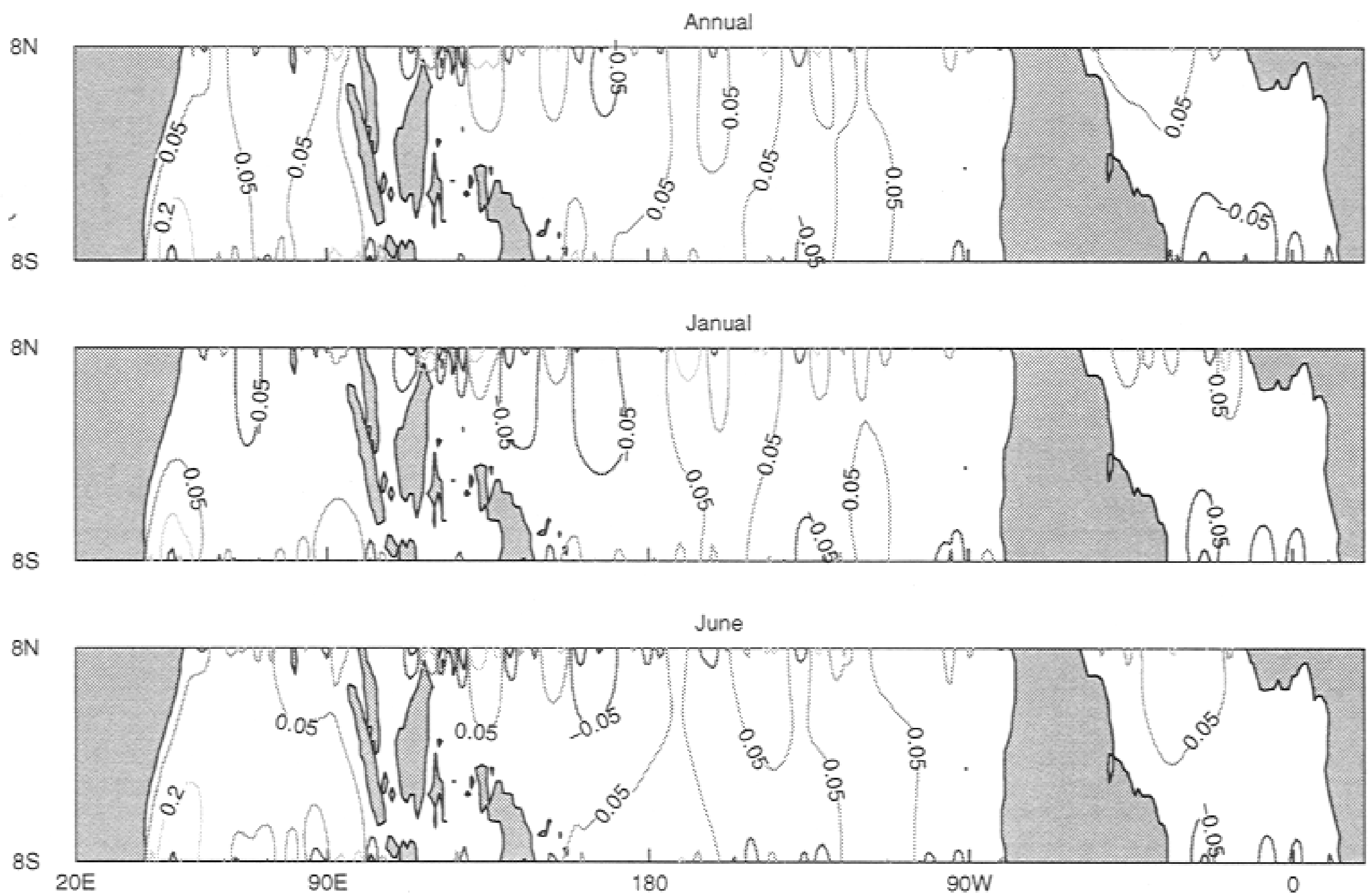
$$\nabla^2 \Pi = \frac{\beta}{A_h} V - \frac{1}{A_h} [\frac{\partial}{\partial x}(\frac{\tau_y}{\rho_0}) - \frac{\partial}{\partial y}(\frac{\tau_x}{\rho_0})].$$

- Boundary Conditions:

  Π Values at 8° S – 8° N

Integration of Munk Model

- Obtaining Π Values in the tropics

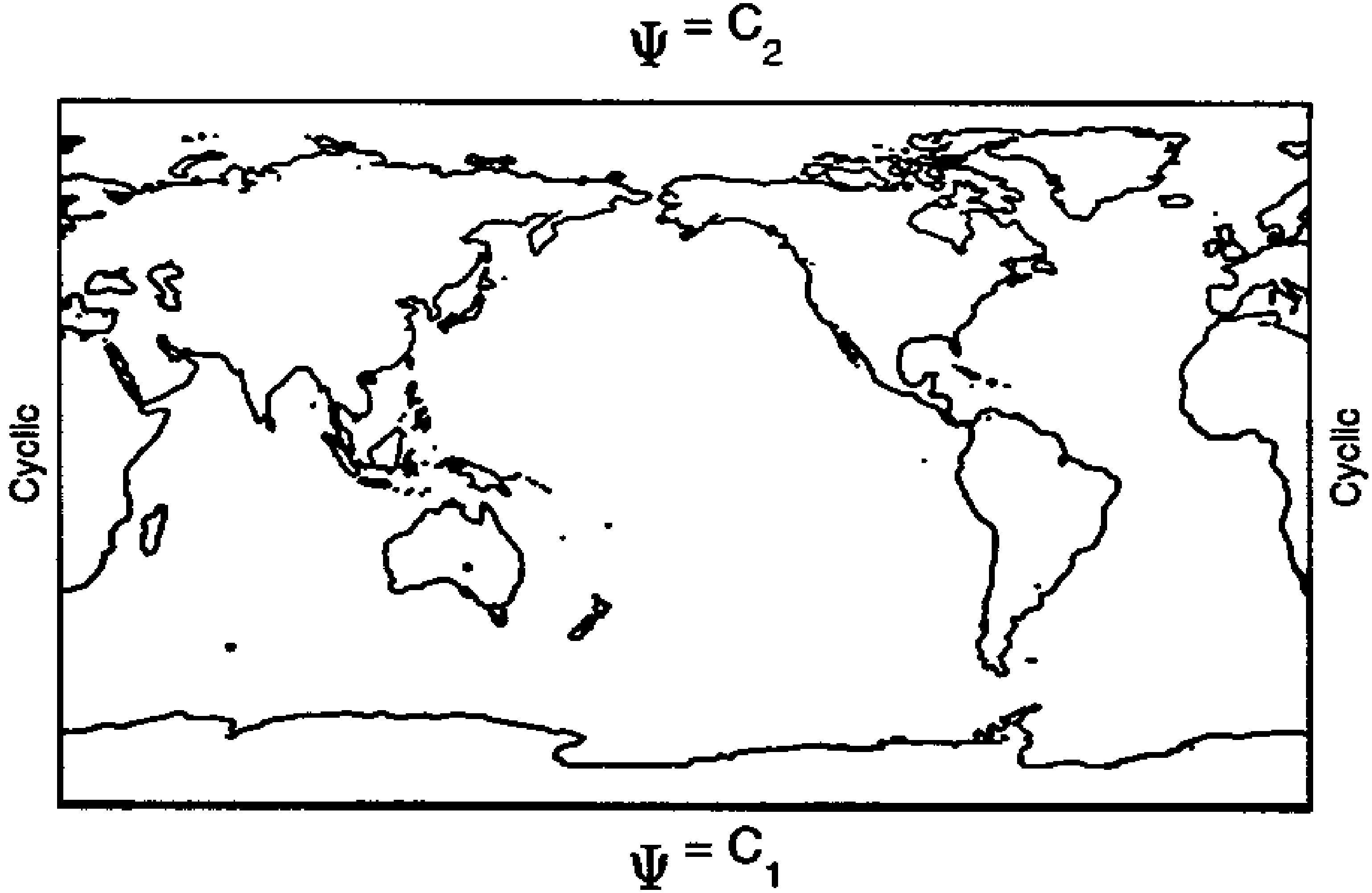


(3) Solving Poisson ψ-Equation

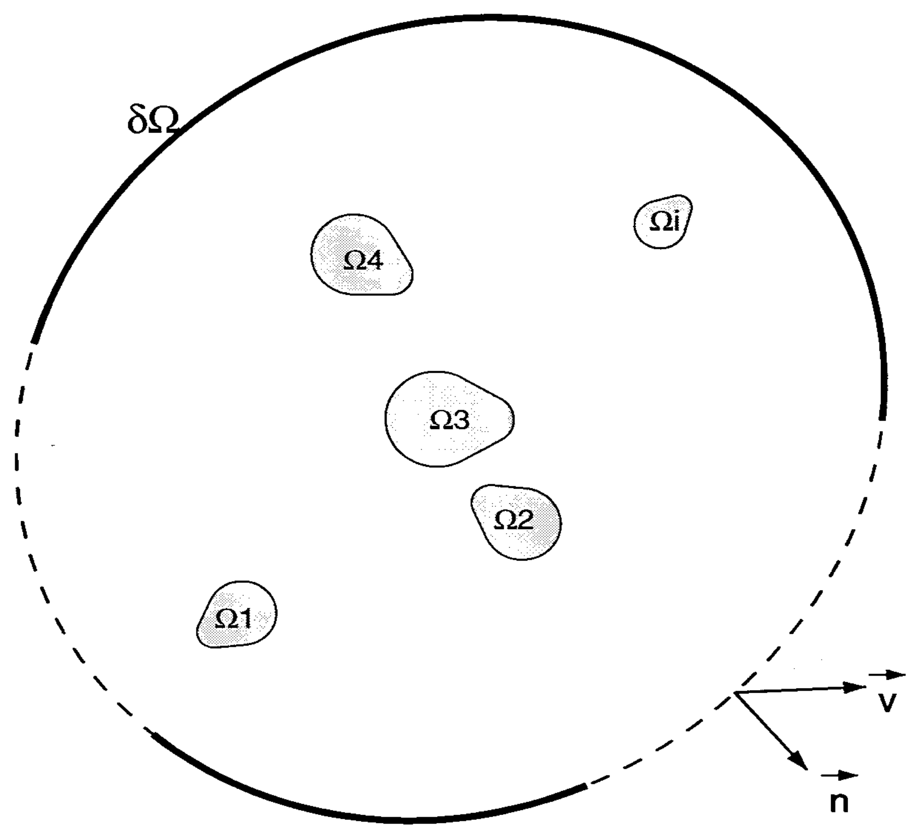
$$\nabla^2 \Psi = \Pi,$$

- With known forcing term for the globe
- We need to know Ψ-Values at islands.

Boundary Conditions for ψ



Multi-Connected Domain



Stokes Theorem

$$-\oint_{\delta\Omega_j} \mathbf{V} \bullet d\mathbf{s} + \oint_{\delta\omega_j} \mathbf{V} \bullet d\mathbf{s} = \iint\limits_{C_j} \mathbf{k} \bullet (\boldsymbol{\nabla} \times \mathbf{V}) dx dy$$

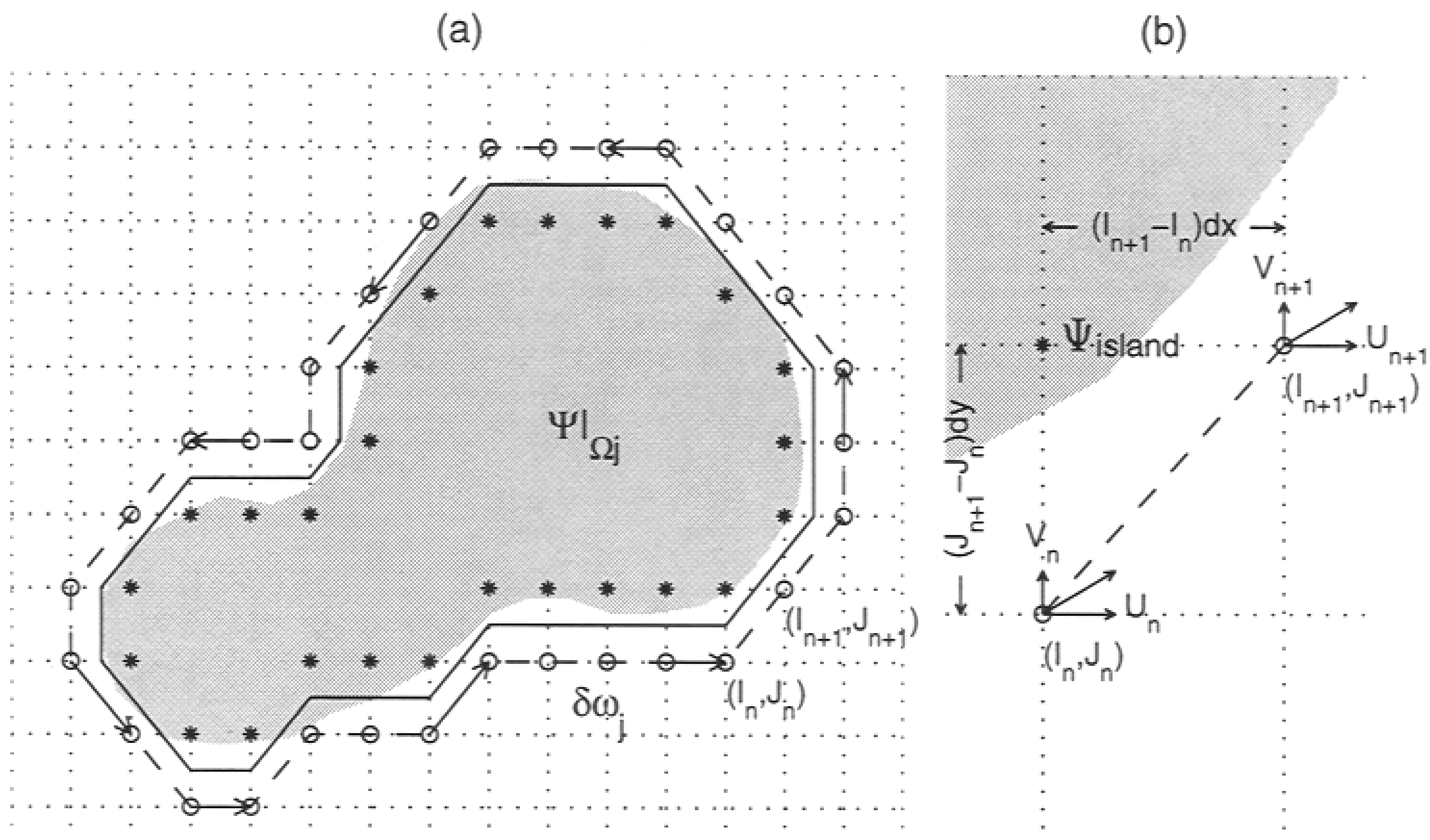
$$\oint_{\delta\Omega_j} \boldsymbol{\nabla}\boldsymbol{\Psi} \bullet \mathbf{n} ds = \oint_{\delta\omega_j} \mathbf{V} \bullet d\mathbf{s} - \iint\limits_{C_j} \mathbf{k} \bullet (\boldsymbol{\nabla} \times \mathbf{V}) dx dy$$

Minimum Circuit Method

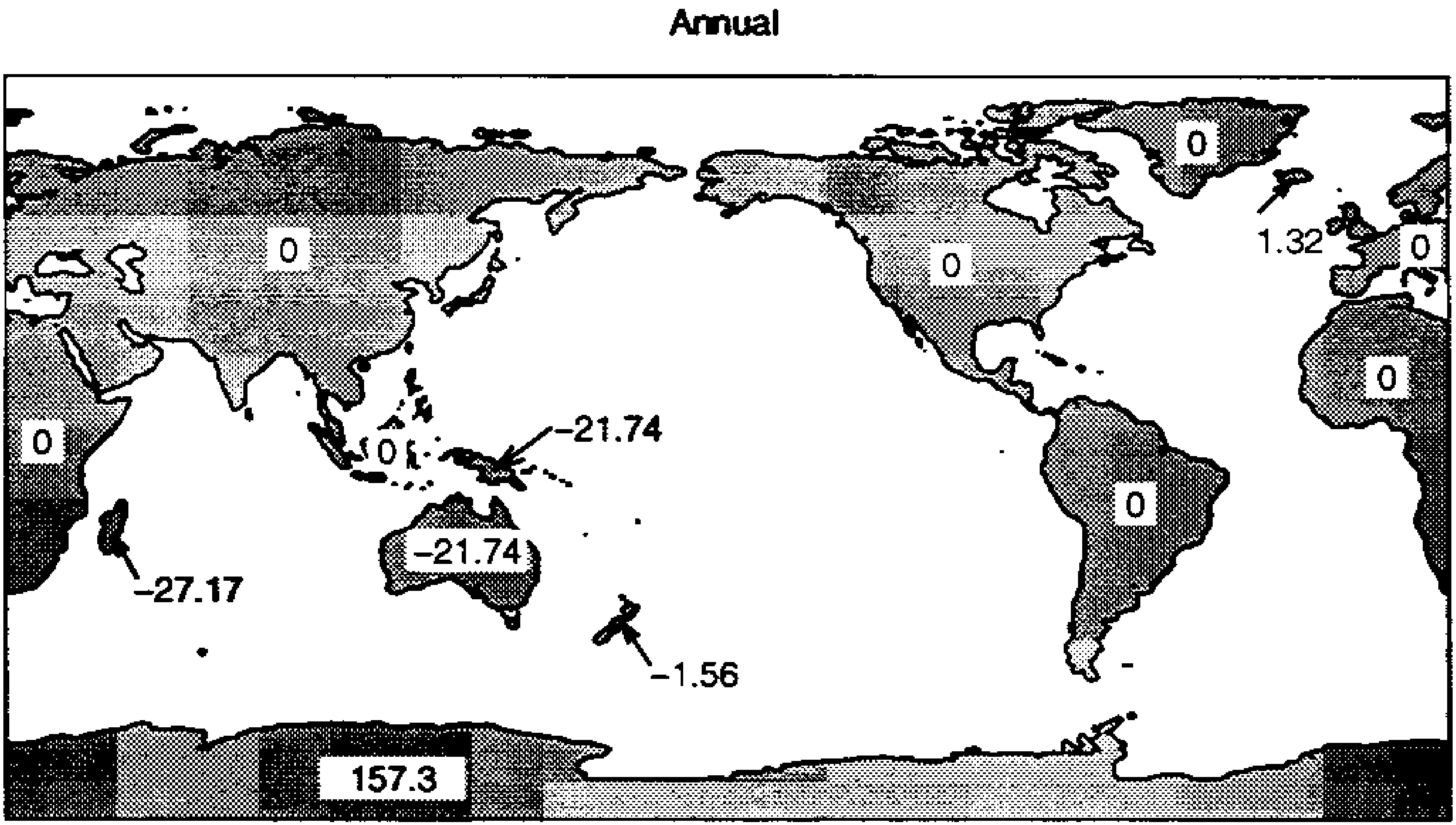
$$\oint_{\delta\Omega_j} \boldsymbol{\nabla}\Psi \bullet \mathbf{n} ds \rightarrow \Gamma_j \quad \text{as } C_j \rightarrow 0$$

$$\Gamma_j \equiv \oint_{\delta\omega_j} \mathbf{V} \bullet d\mathbf{s}$$

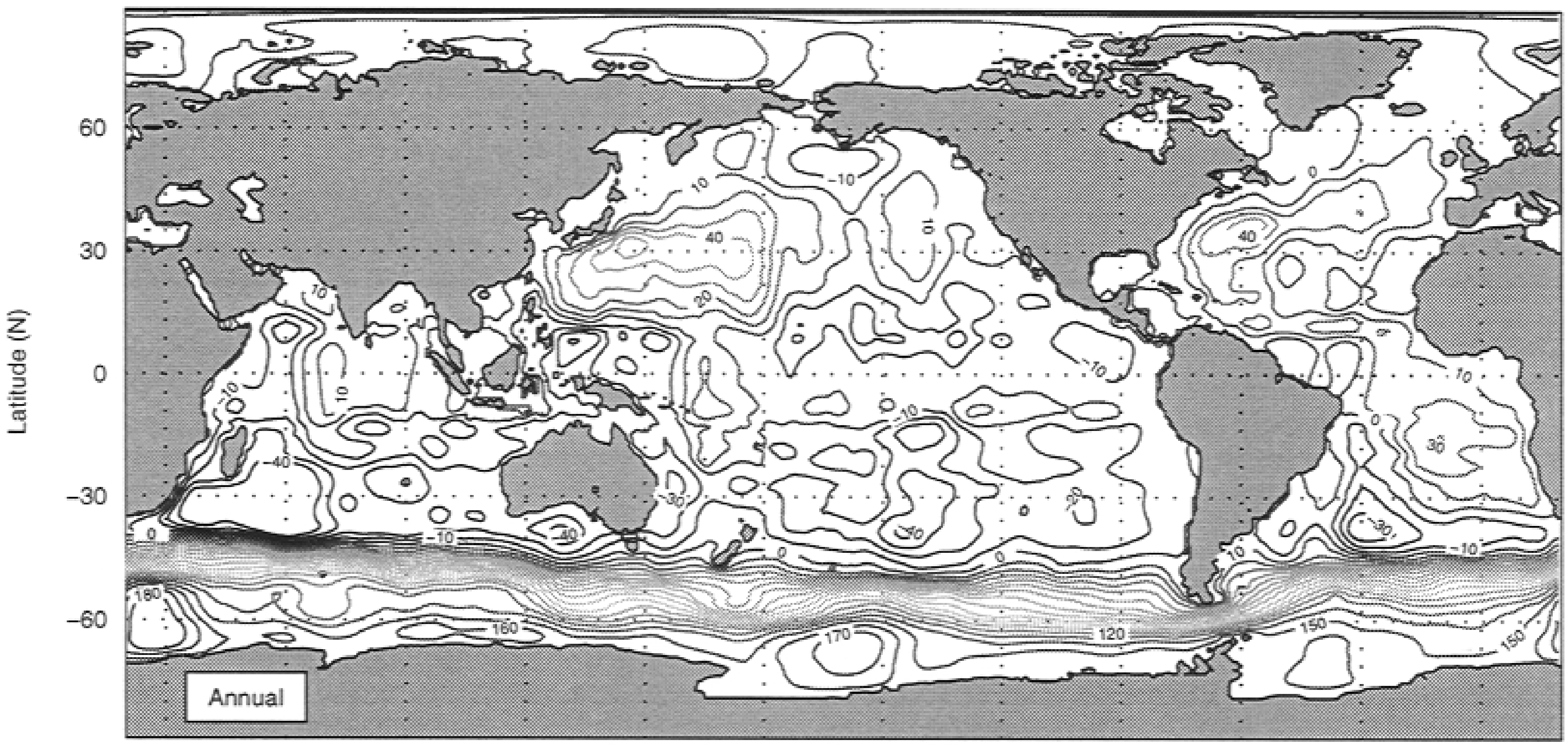
Stokes Theorem



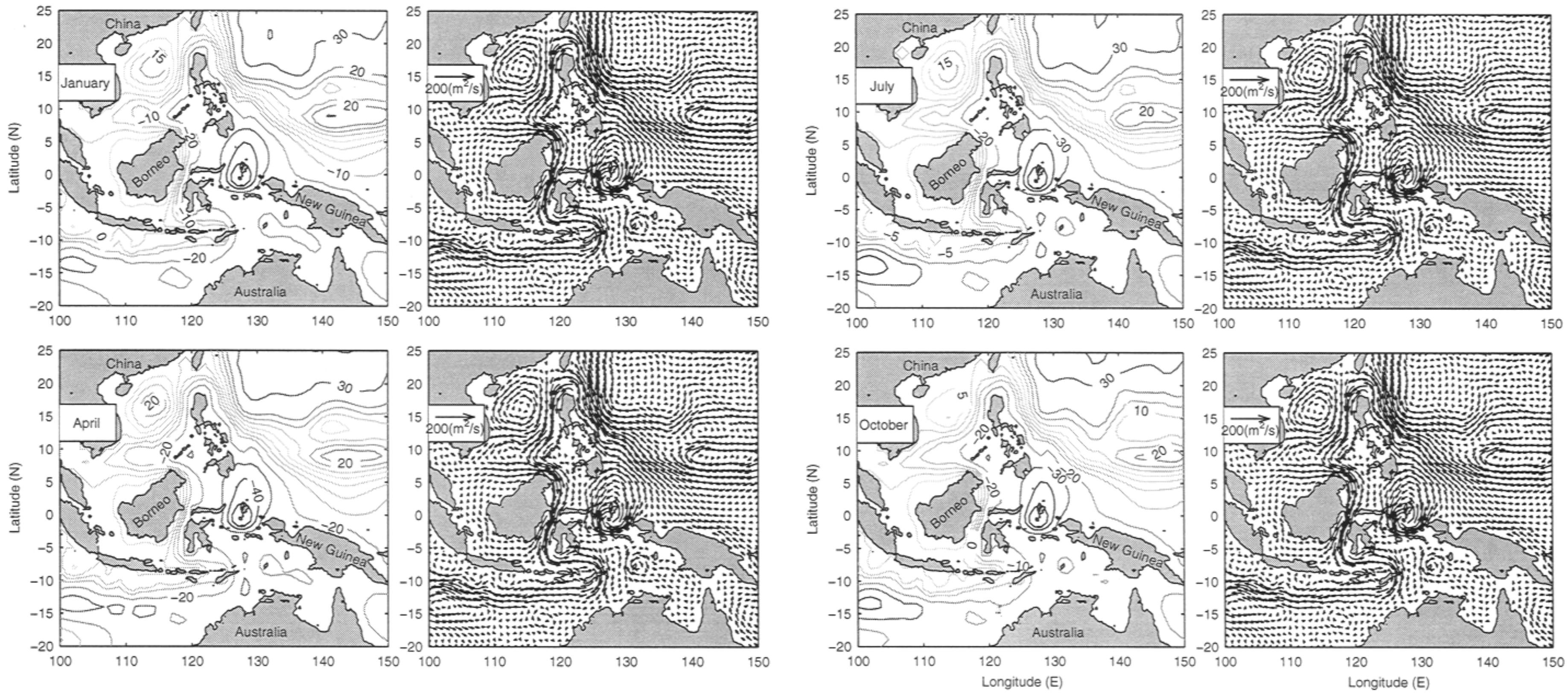
Ψ-Values at Islands (Annual Mean)



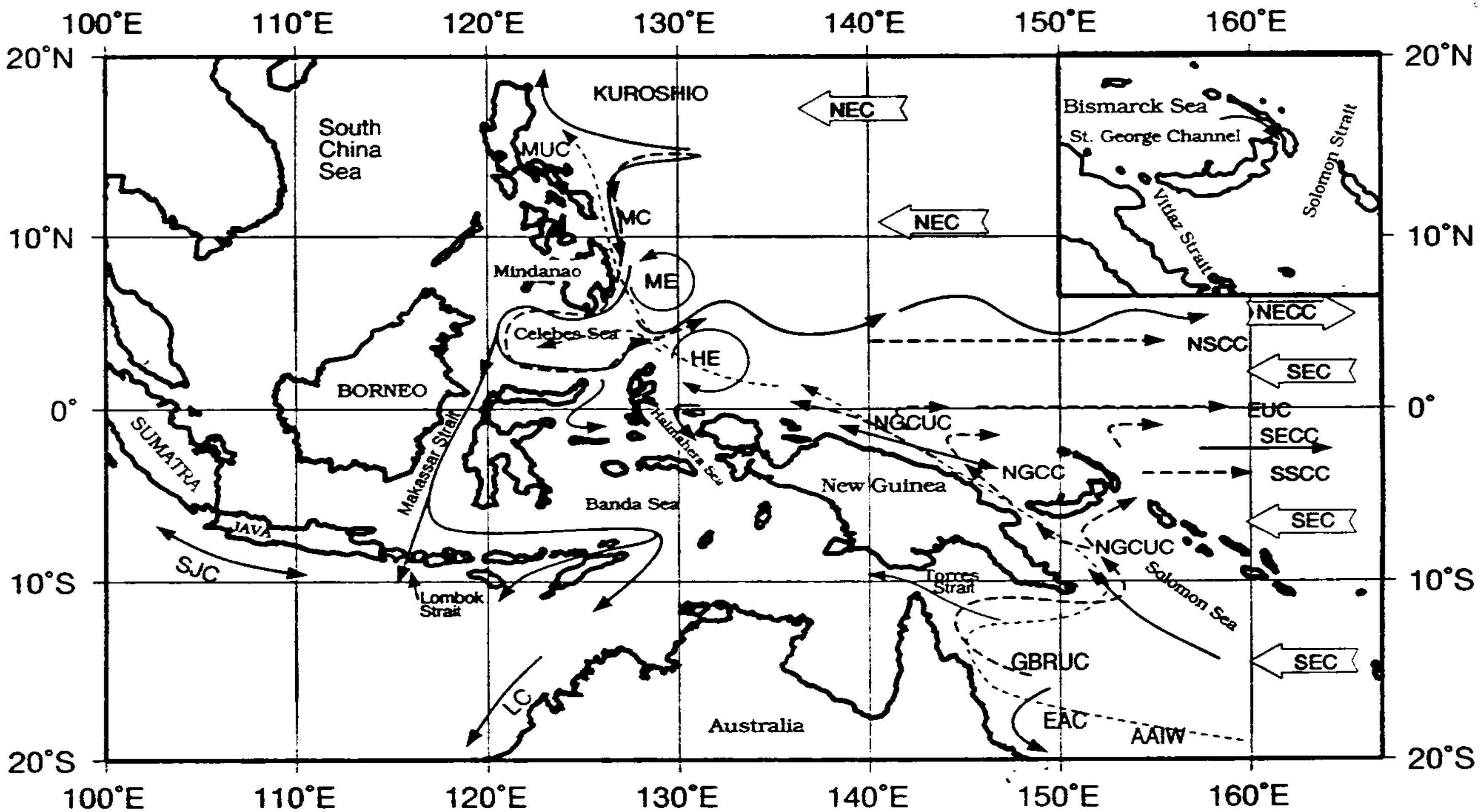
Global Volume Transport Streamfunction



Indonesia Throughflow
(Pacific – Indian Link)



Indonesia Throughflow (Fine et al. 1994)



Australia-Bali Section Transports (Fieux et al. 1994)

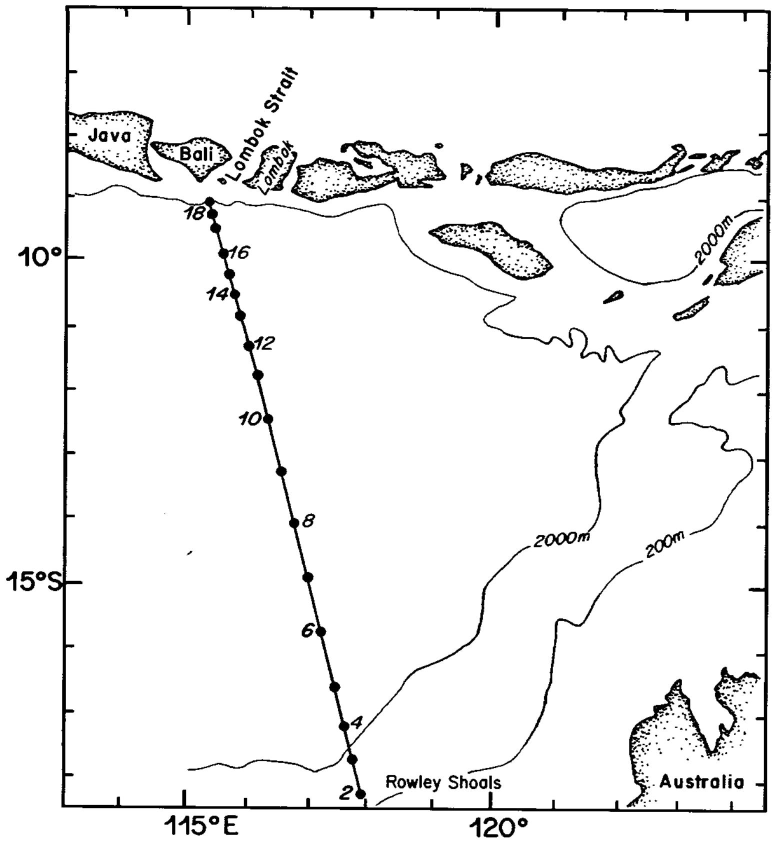
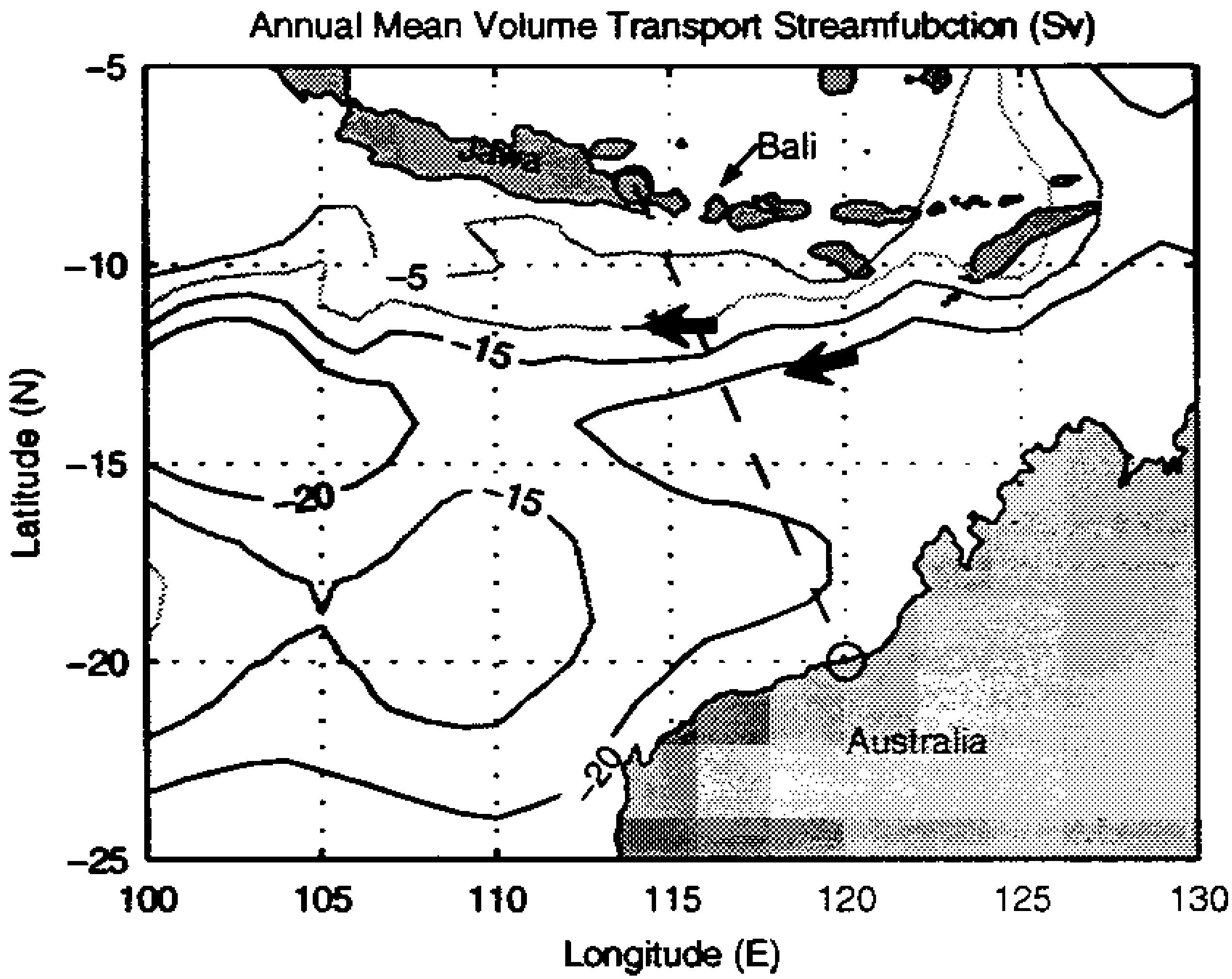


Table II-14: Australia–Bali Section Transports (in Sverdrups; adapted from Fieux *et al.*, 1994)

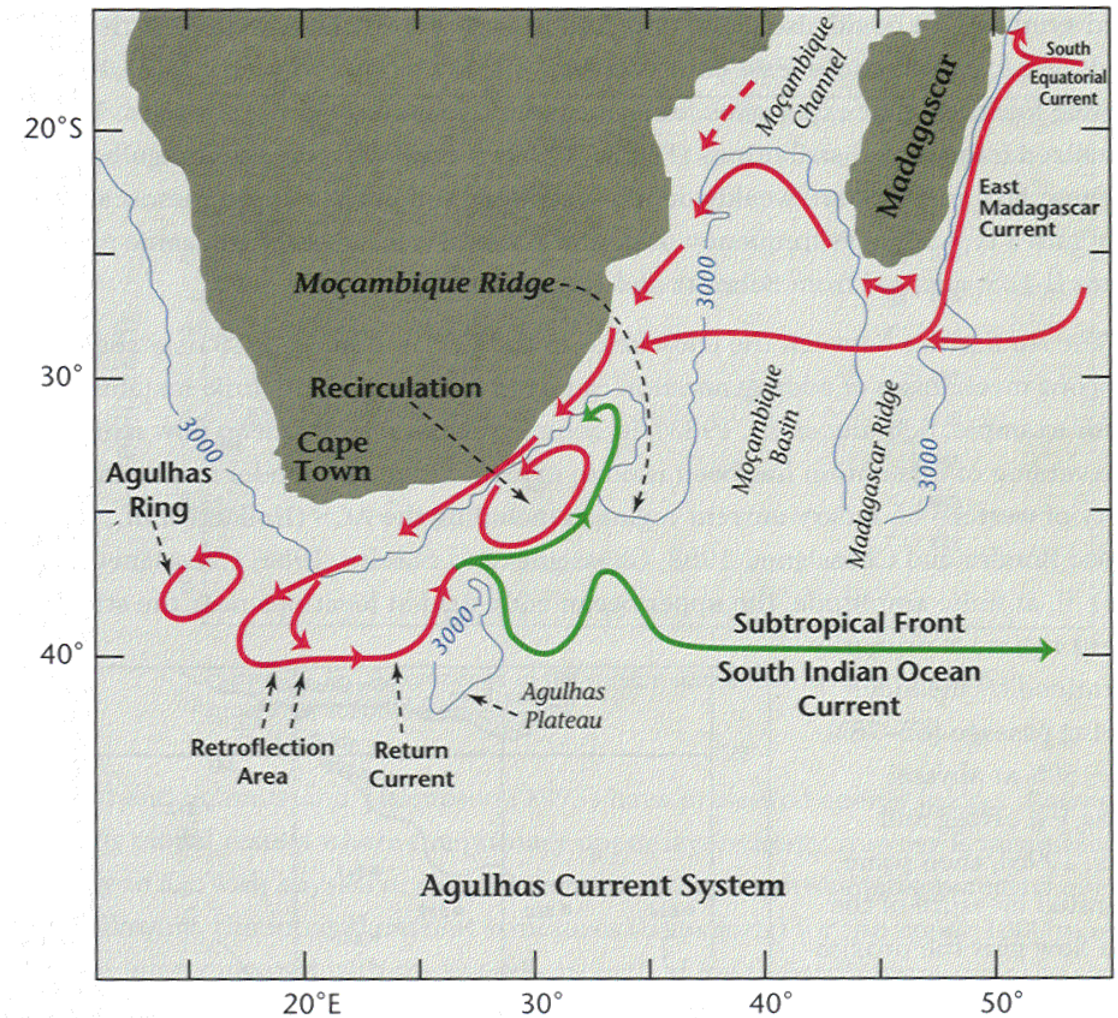
Layers	Transport*
0–200 db	–23.1
200–500 db	–2.7
Total 0–500 db	–25.8
500–2000 db	+9.6
Total 0–2000 db	–16.2

*Minus signs indicate westward flow.

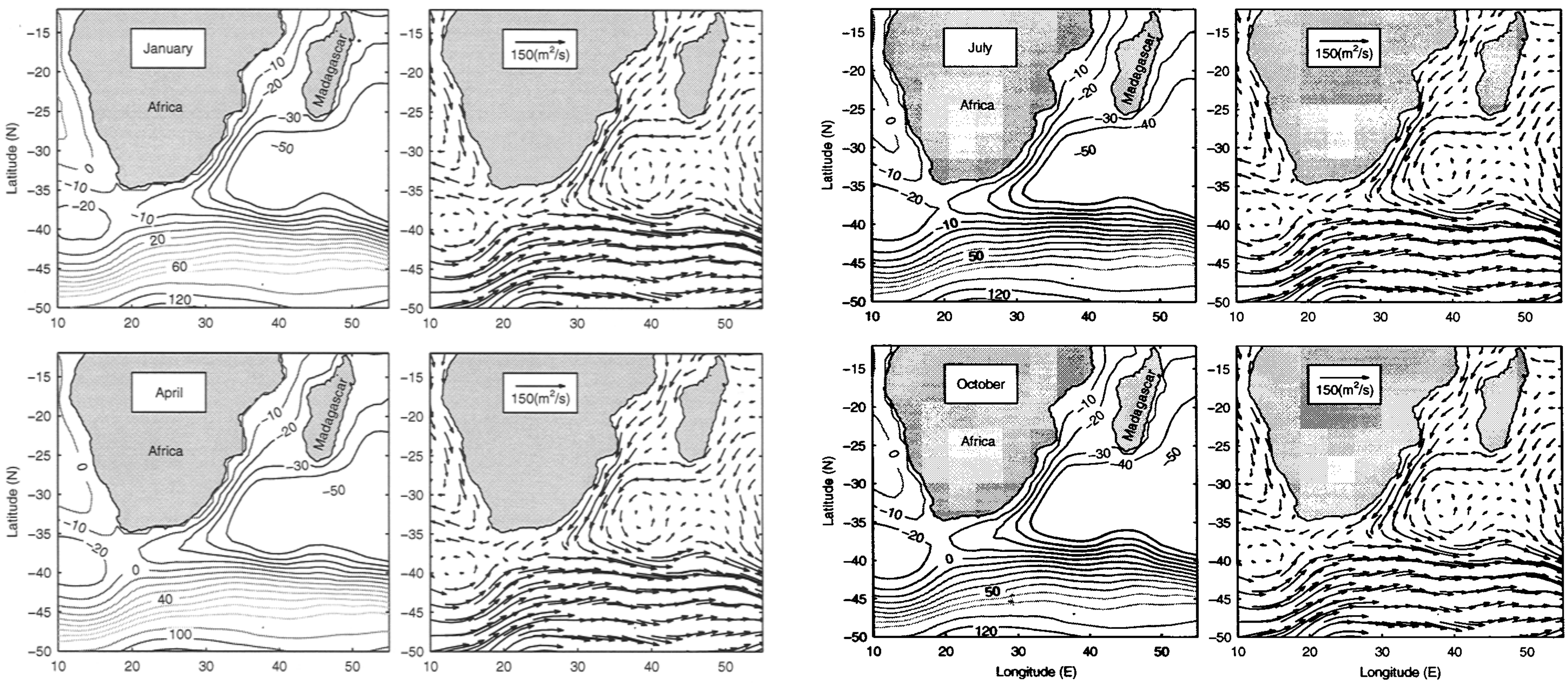
Annual Mean Transports Calculated Using Ekman-Munk Model



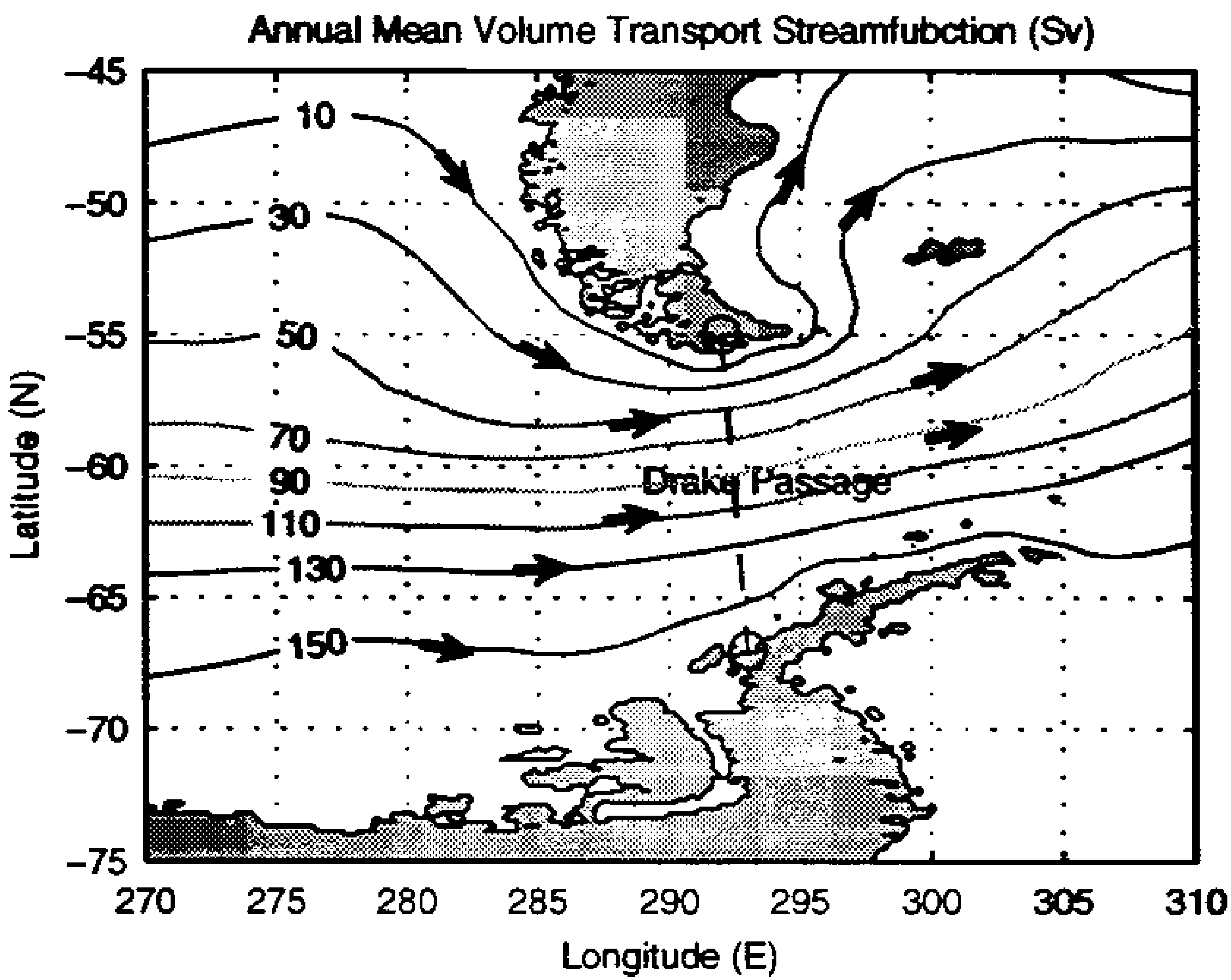
South Atlantic Link to the Indian Ocean: Agulhas Current System (Schmitz 1996)



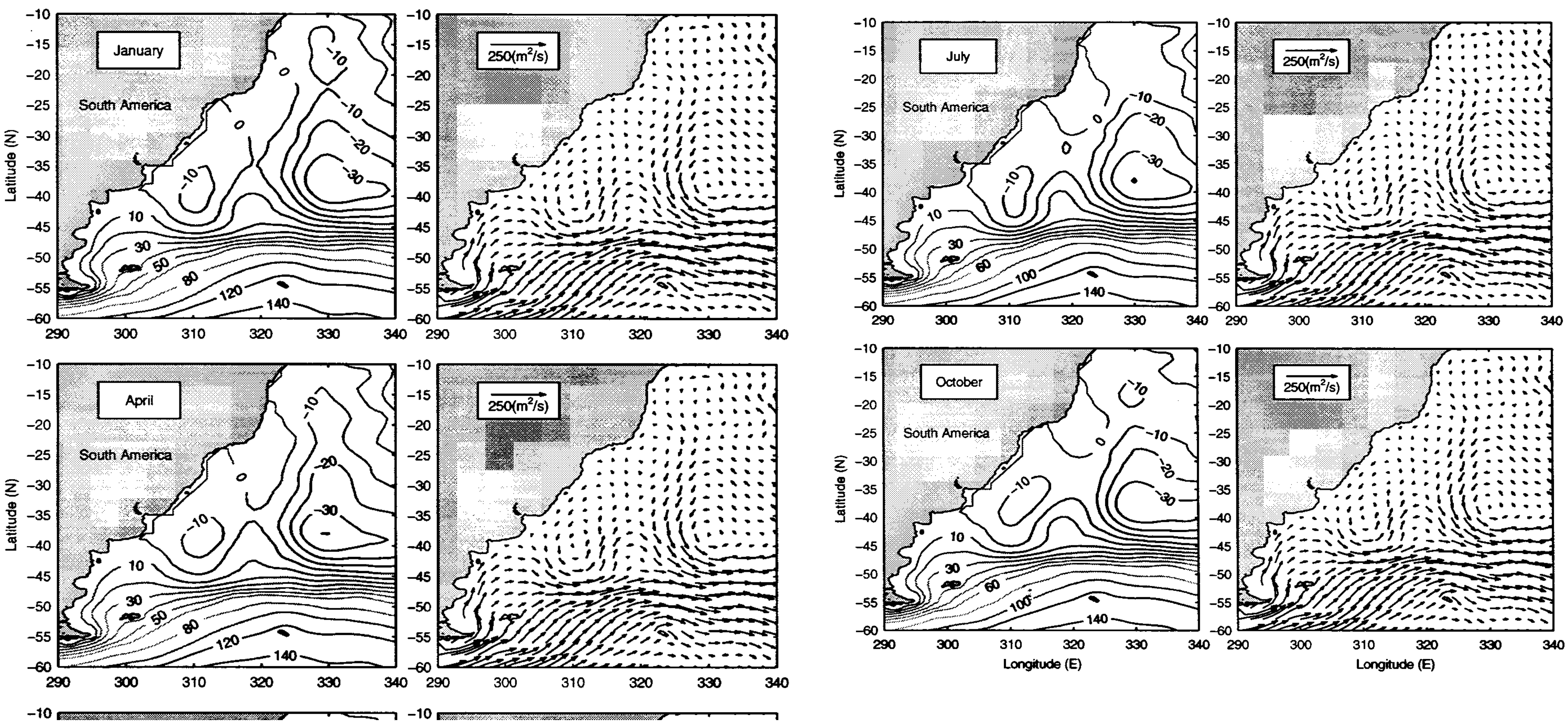
Agulhas Current System
(from the Ekman-Munk Model)



Atlantic Linkage to Pacific Ocean



Malvinas Confluence



Conclusions

- Ekman-Munk model has capability to diagnose the volume transport from wind and hydrographic data
- Stokes theorem is effective for determining streamfunction at islands
- Annual and monthly mean global volume transport data are useful for coastal modeling
- The adaptation package is easily incorporated into any coastal model